

Abstract

This paper presents a detailed analysis of the September 8-9 2002 flash flood event in the Gard region (southern France) using two distributed hydrological models: CVN built within the LIQUID[®] hydrological platform and MARINE. The models differ in terms of spatial discretization, infiltration and water redistribution representation, and river flow transfer. MARINE can also account for sub-surface lateral flow. Both models are set up using the same available information, namely a DEM and a pedology map. They are forced with high resolution radar rainfall data over a set of 18 sub-catchments ranging from 2.5 to 99 km² and are run without calibration. To begin with, models simulations are assessed against post field estimates of the time of peak and the maximum peak discharge showing a fair agreement for both models. The results are then discussed in terms of flow dynamics, runoff coefficients and soil saturation dynamics. The contribution of the sub-surface lateral flow is also quantified using the MARINE model. This analysis highlights that rainfall remains the first controlling factor of flash flood dynamics. High rainfall peak intensities are very influential of the maximum peak discharge for both models, but especially for the CVN model which has a simplified overland flow transfer. The river bed roughness also influences the peak intensity and time. Soil spatial representation is shown to have a significant role on runoff coefficients and on the spatial variability of saturation dynamics. Simulated soil saturation is found to be strongly related with soil depth and initial storage deficit maps, due to a full saturation of most of the area at the end of the event. When activated, the signature of sub-surface lateral flow is also visible in the spatial patterns of soil saturation with higher values concentrating along the river network. However, the data currently available do not allow the assessment of both patterns. The paper concludes with a set of recommendations for enhancing field observations in order to progress in process understanding and gather a larger set of data to improve the realism of distributed models.

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2 **Keywords:** Flash floods, distributed modelling, soil variability, hydrological processes,

3 saturation excess, infiltration excess, model inter-comparison

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1 Introduction

Flash floods represent the most destructive natural hazard in the Mediterranean region, causing around one billion Euros worth of damage in France over the last two decades (Gaume et al., 2009). Flash floods are associated with extreme and rare rainfall events and usually occur in ungauged river basins. Amongst them, small-ungauged catchments are recognized as the most vulnerable to storm driven flash floods (Ruin et al., 2008).

In order to limit damage to the population, there are several currently accepted methods for predicting flash floods in ungauged river basins. The flash flood guidance (Georgakakos, 2006; Norbiato et al., 2008) and the discharge threshold exceedance approach (Reed et al., 2007; Younis et al., 2008) are built to give an early flash flood warning with suitable time to organize the civil protection. These operational methods are very efficient for warning, but must be complemented with field experiments and modelling studies to improve the understanding of the major hydrological factors associated with the flood events. In order to progress in process understanding related to flash floods, large scale in situ experiments are scheduled in the context of the HyMeX program (www.hymex.org). This project aims at improving our understanding and prediction of the Mediterranean Sea water balance. The latter can be highly impacted by extreme events which provide a sudden input of fresh water to the sea (Drobinski et al., 2008). One of the focuses of HyMeX is therefore extreme events and, in particular, flash floods over the whole Mediterranean region. The objective is to better understand these events and to improve the predictive capability of hydro-meteorological models in simulating and anticipating them. The quantification of global change impact on the frequency and magnitude of these extreme events will also be analyzed within HyMeX. To achieve these goals, the HyMeX program is planning to enhance the observation capabilities of the scientific community in the Mediterranean region in order to better document extreme rainfall events and flash floods during a ten-year period, with an enhanced period of four

1 years and two special observation periods in the fall. This experimental framework offers a
2 good opportunity to enhance hydrological process observations and understanding. Progress
3 in hydrological modelling of these events are also expected. As all the catchments cannot be
4 surveyed, it is important to determine which type of observations are required and, where and
5 when these observations are needed.

6 Distributed hydrological models, representing physical mechanisms, are interesting tools for
7 hypothesis testing and field design (Loague et al., 2006). Sensitivity studies to process
8 representation, spatial discretization, input data and parameters can be performed that allow
9 the quantification of the impact of various functioning hypotheses on the hydrological
10 response (Pinol et al., 1997, Sangati et al., 2009). Vivoni et al. (2007), using the tRIBs
11 distributed model (Ivanov et al., 2004), explored the complex interactions between the various
12 runoff contributions (infiltration excess, saturation excess, perched return flow, groundwater
13 exfiltration) and the rainfall and catchments characteristics (soil, land use, topography). They
14 showed how various responses can be observed at the outlet according to spatial and temporal
15 variability of these factors and that threshold effects can be observed. Using the same model,
16 Noto et al. (2008) focused on the impact of initial moisture (specified using a variable initial
17 groundwater level) on the catchment response; they highlighted the complexity of the
18 hydrological response to rainfall and soil characteristics. These studies focused on synthetic
19 rainfall events and the studied catchment was about 800 km². Extreme events were not
20 considered in these studies and the analysis was performed for the whole catchment without
21 describing the internal variability.

22 The present paper deals with extreme events at the regional scale and aims at addressing small
23 ungauged catchments ranging from a few km² to about 100 km². In the context of the PUB
24 (Prediction of Ungauged Catchments) initiative, the questions addressed in the paper are the
25 following: i) is it possible to set up physically-based distributed hydrological models at the

1 regional scale using available data and information for flash flood simulation?; ii) are post-
2 flood data of maximum peak discharge useful to assess the relevance of the modelling?; iii)
3 are they relevant to discriminate between various model structures?; iv) are sensitivity studies,
4 based on distributed hydrological models outputs, useful to assess the limits of current
5 observations and highlight which information should be acquired in future field experiments,
6 in order to progress in the understanding and simulation of flash flood events for such
7 catchments. For this purpose, two distributed hydrological models with different model
8 structures are used. The study is conducted for the 8-9th September 2002 event, which affected
9 the Gard region in south-east France. This event was exceptional, both in its extent (more than
10 20000 km² affected) and duration, with more than 600 mm accumulated rainfall in 24 h in
11 some locations. For this event, radar rainfall and post flood field data of maximum peak
12 discharge are available.

13 The case study, rationale for model choice, model description and set up are presented first.
14 The methodology used for model evaluation based on maximum peak discharge and
15 sensitivity studies is also presented. The second part of the paper describes the model results
16 in terms of simulation of maximum peak discharge. This analysis is complemented by local
17 sensitivity studies and a discussion of model results in terms of hydrograph, runoff
18 coefficient, and soil saturation dynamics. After a discussion of the results, guidelines for
19 future experiments are proposed. These includes the processes, variables and parameters that
20 require further description and investigation.

Materials and methods

Case study and available data

The case study is the September 8-9, 2002 event which affected the Gard region located in the South-East of France (Figure 1). This event is the most important event ever recorded in this region. It was responsible for 24 casualties and caused roughly 1.2 million euros worth of damage. This event was extensively described in Delrieu et al. (2005) and in Manus et al. (2009). Thus, we shall only present the data used in the present study. In terms of rainfall input, we used rainfall intensity data from the Bollène radar with a $1 \times 1 \text{ km}^2$ grid resolution and a 5min time step, with the ST-AD3 processing protocol described by Delrieu et al. (2009). The spatial variability of soils is described using the Languedoc-Roussillon soil data base (later referred as BDSol-LR), provided by INRA (National Institute of Agronomic Research) from the French IGCS (Inventory, Management and Conservation of Soils) program. This database provides information (i.e. texture, horizon depth, etc..) on pedological landscape units called Soil Cartographic Units (SCUs). These units are established with a resolution of 1/250000 and they are geo-referenced. They are composed of Soil Typological Units (STUs), the vertical heterogeneity of which is described by stratified homogeneous layers of soil. The proportion of STUs is given within a particular SCU, but the precise location of STUs within this SCU is unknown. Each STU is described through tables providing both quantitative and qualitative information from which quantities such as percentage of sand, clay, silt, organic matter or soil depth can be derived. Pedo-transfer functions are used to derive the hydraulic parameters of the various soil horizons (see details below).

For model evaluation, we use data from an extensive post-flood investigation carried out during the months following the event. The methodology of Gaume and Bouvier (2004) was used during this field survey. The survey gathered a regional information about the flood, allowing the analysis of the hydrological behaviour of watersheds with an area of 2 to 300

km². The procedure provides estimation of maximum discharges based on water level marks and simple hydraulic hypotheses for the derivation of the flow velocity. The flood chronology is documented based on witnesses interviews. A sub-set of data corresponding to the same 17 sub-catchments, already studied by Manus et al. (2009), is chosen. The area of these catchments ranges from 2.5 to 50 km² (Table 1). In addition, one 99 km² gauged catchment, the Saumane catchment, is also considered. Catchment locations are shown in Figure 1 and their main characteristics are summarized in Table 1. The catchments located in the north-western part of the domain have steep slopes, whereas the catchments located in the south-eastern part of the domain are situated in flatter areas (Figure 1). Table 1 shows a large variability in average soil depth and maximum storage capacity of these catchments. The variables in Table 1 are derived from the BDsol-LR using the dominant STU in each SCU.

Modelling hypotheses and model choice

First of all, the choice of the model should be dictated by the objectives of the modelling exercise (e.g. Kamp and Burges, 2007). Concerning regional flash flood modelling, model requirements are presented in Borga et al. (2008) and Sangati et al. (2009). They underline that “*the space-time excess rainfall distribution and drainage network structure provide the most important control on extreme flood response structure*” (Sangati et al., 2009). Therefore, models used in regional flash flood studies should take into account high resolution rainfall and represent soil spatial variability. In order to conduct the study, two distributed hydrological models, able to cope with these requirements, and differing in their structure are chosen. The first model is the CVN (Cévennes) model which is built within the LIQUID[®] hydrological modelling platform (Viallet et al., 2006; Branger et al., 2009¹). The second

¹ Branger, F., Braud, I., Debionne, S., Viallet, P., Dehotin, J., Hénine, H., Nédélec, Y., Anquetin, S., 2009. Towards multi-scale integrated hydrological models using the LIQUID framework. Overview of the concepts and first application examples, *Environmental Modeling & Software*, in revision

model is the MARINE model, dedicated to flash flood modelling (Maubourguet et al., 2007; Bessière et al., 2008). The two models are representative of the variety of distributed hydrological models used in flash flood event simulation (see references cited by Sangati et al., 2009). In addition and in order to be useful for experimental design, the models must be physically-based so that their inputs parameters and variables, and outputs variables can be related to observable quantities. Kampf and Burges (2007) propose a synthesis and a classification of distributed hydrological models, that highlights the main points which must be considered when describing a model structure. Following this classification, the spatial domain over which the model is applied, the space and time discretization, and the processes which are considered, including surface/subsurface, overland, channel flows, and other processes, as well as their coupling, must be described. These various points are reviewed below for the two models, following the suggestion of Kampf and Burges (2007). Both models are applied at the catchment scale.

Models description

MARINE

The MARINE model is a spatially distributed rainfall-runoff model dedicated to extreme event simulation and developed on the basis of physical process representation. The model is structured as three physical modules (Figure 2a), which represent vertical transfer (infiltration and runoff generation for the soil component), lateral transfer and both the overland flow component and the flow through the drainage network. Interception and evapotranspiration are not represented since they are considered negligible during individual flood events.

The model output is a simulated hydrograph which is available at the catchment outlet but also at any point of the drainage network. It is also possible to follow the evolution of

distributed variables such as soil moisture over all the catchment. A description of each model component is detailed hereafter.

Spatial discretization

The spatial discretization of the catchment is performed using the Digital Elevation Model grid resolution.

Surface and subsurface flow

Infiltration is described using the Green and Ampt (1911) model, which assumes one-dimensional flow in vertically homogeneous soil columns.

The subsurface model is based on Darcy's law. Using the classical assumptions that (i) the slope of the water table in the saturated zone coincides with the local topographic slope, and (ii) the local transmissivity is an exponential function of the local storage deficit (Original TOPMODEL assumption (Beven and Kirkby, 1979)), the flow per unit width q is expressed as:

$$q(t) = T_0 \exp\left(-\frac{\theta_s - \theta}{m}\right) \tan \beta \quad (1)$$

where T_0 is the local transmissivity of fully saturated soil ($\text{m}^2.\text{s}^{-1}$), θ_s and θ are the saturated and local water contents ($\text{m}^3 \text{ m}^{-3}$), m is the transmissivity decay parameter (-) and β is the local slope angle (-). Soil water can exfiltrate in two cases: (i) when the soil water content, θ , exceeds saturated water content θ_s ; (ii) when the soil water reaches the drainage network. It is assumed that exfiltration into the drainage network occurs with a velocity calculated using equation (1). Both infiltration excess and saturation excess are therefore represented within MARINE. In the present study, and in order to allow comparison with the CVN model, the MARINE model is run without activation of the subsurface lateral flow component. In the sensitivity analysis, the impact of including subsurface flow in the modelling is discussed.

Overland flow and channel flow

The surface runoff is divided into two parts: the overland flow and the flow through the drainage network. Both are simulated using the 1D kinematic wave approximation of the Saint-Venant equations with the Manning friction law. When the drainage area becomes greater than 1 km², the overland flow is structured in a drainage network. Flow in this drainage network takes into account a simplified network reach cross-section (Figure 2b). This leads to a new transport equation in the drainage network. Characteristics of the network reach are calculated using geomorphological considerations (Liu and Todini, 2002): network reach width W_{Di} and depth H_{Di} are taken to increase as a function of the area drained by the i^{th} cell, such that :

$$\begin{aligned} W_{Di} &= W_{Dmax} + \left(\frac{W_{Dmax} - W_{Dmin}}{\sqrt{a_{tot}} - \sqrt{a_{th}}} \right) (\sqrt{a_{di}} - \sqrt{a_{tot}}) \\ H_{Di} &= H_{Dmax} + \left(\frac{H_{Dmax} - H_{Dmin}}{\sqrt{a_{tot}} - \sqrt{a_{th}}} \right) (\sqrt{a_{di}} - \sqrt{a_{tot}}) \end{aligned} \quad (2)$$

where W_{Dmax} is the maximum width at the basin outlet, W_{Dmin} is the minimum width, corresponding to the threshold area $a_{th}=1$ km², which is the minimum upstream drainage area required to initiate a channel, a_{tot} is the total area and a_{di} is the area drained by the i^{th} cell.

MARINE is run with a fixed time step of 15 s for numerical stability reasons.

CVN

The CVN model is built within the LIQUID[®] modelling platform (Viallet et al., 2006). The system allows a customized model building by assembling process modules described with their own spatial discretization and numerical schemes. These modules can exchange fluxes both in space and time, and the interactions between the various modules are managed by a time sequencer. Each model component is therefore run with an adaptative time step, consistent with its own dynamics. Using the LIQUID[®] modelling platform, model complexity can be increased step by step, allowing a better control and analysis of the impact of various

1 model conceptualizations in terms of hydrological response. In this study, a first –simple-
2 version of the model is used. In particular, as in the MARINE model, interception and
3 evapotranspiration are neglected during the event.

4 ***Spatial discretization***

5 The catchment discretization is based on the principles proposed by Dehotin and Braud
6 (2008) with several levels of discretization. The first level is composed of sub-catchments,
7 organised along the river network (Figure 3). The second level allows the introduction of
8 some spatial heterogeneity within the sub-catchments, and leads to hydro-landscapes that are
9 assumed homogeneous in terms of hydrological response. Rainfall input is also averaged at
10 the scale of the hydro-landscapes. For flash-flood modelling, only soil heterogeneity as
11 described by the pedology map is considered in the hydro-landscape delineation. The result is
12 illustrated in Figure 3. The hydro-landscapes are the elementary units for the modelling of soil
13 infiltration and vertical water redistribution. The hydro-landscapes are vertically discretized
14 into soil horizons with different hydraulic properties.

15 ***Surface and subsurface flow***

16 On each hydro-landscape, the Richards equation is solved using the Ross (2003) method as
17 validated more extensively by Varado et al. (2006) and Crévoisier et al. (2009). The horizons
18 are sub-divided into cells that are 1cm thick for the stability of the numerical scheme. When
19 the topsoil is saturated, ponding is generated. At the bottom of the hydro-landscapes,
20 gravitational or zero flux boundary conditions can be imposed. Subsurface lateral transfer
21 between hydro-landscapes is not considered in this first version of the model.

22 ***Overland and channel flow***

23 Transfer within the river network is performed using the 1D kinematic wave approximation of
24 the St-Venant equation. The network is discretized into river reaches with one reach per sub-
25 catchment (Figure 3). The river section is assumed trapezoidal with geometrical parameters

assigned according to the Strahler order. At this development stage of the model, the routing scheme allowing the transfer of ponding generated on the hydro-landscapes to the river reaches is very simple, as we are focusing on peak discharge and risk assessment,. The ponding is quasi-instantaneously transferred to the closest river reach. The implications of these simplifying hypotheses have been discussed by Manus et al. (2009) (see interactive discussion) and are thus not repeated here.

Model set up

In order to be able to quantify differences and similarities in hydrological responses between the two models, the same input data were used whenever possible. No calibration is performed and both models are run using the a priori information and expert knowledge on the various models derived from previous studies or model user expertise. These simulations will be referred to as the reference simulations in the following of the paper.

Spatial discretization

Both models rely on Digital Elevation Model (DEM) data for river network and sub-catchments delineation. The same 75 m resolution DEM, provided by IGN is used.

For the CVN model, the D8 algorithm (e.g. Tarboton, 1997) is used to define the drainage direction map, the cumulated drainage area and the river network. A minimum drainage area of 0.1 km^2 is used to distinguish between network and hillslope pixels. The Strahler order one sub-catchments are considered for the definition of the first level of discretization (Figure 3). This ensures that the hypothesis of instantaneous transfer of ponding to the closest river reach is acceptable. For each river reach, the average slope is computed from the DEM slope information. Unfortunately, the DEM information is truncated to 1 m in height, which can sometimes lead to zero or very small slopes. This is corrected by assigning a minimum slope of 0.5 m (half the vertical DEM resolution) divided by the river reach length.

This problem of DEM resolution is much more problematic for the MARINE model which requires a continuity of flow at the scale of the DEM grid. To avoid problems, the DEM is used with a 150 m resolution in MARINE. It is processed to identify connections between cells, leading to the determination of the catchment extent and the flow pathways. DEM data also allow the retrieval of hillslopes and cumulated drainage area, the drainage network and the geometric characteristics of the network reaches. Drainage network extreme widths and depths are derived from in situ measures. As in many rainfall-runoff models (Liu and Todini, 2002), drainage is only possible to the North, East, South or West for the four adjacent cells at each edge in MARINE (D4 algorithm).

Figure 4b shows the 30 h accumulated rainfall over the catchments. Rainfall is distributed at the pixel scale within MARINE and at the hydro-landscape scale within CVN.

Surface and subsurface flow

For the specification of soil characteristics, the same pedology map, from the BDsol-LR is used by both models. As mentioned above, several Soil Typological Units (STUs) can be present within the Soil Cartographic Units (SCUs). In the present study, the dominant STU within each SCU was assumed representative of the whole SCU. A common map of soil depth and soil texture was built using this information.

Within MARINE, soils are assumed to be vertically homogeneous. Soil classes are assigned according to the Rawls and Brakensieck (1983) soil classes according to their texture. For each soil class, a set of parameters (saturated water content, soil suction for the Green and Ampt model and saturated hydraulic conductivity) is assigned.

Within the CVN model, the vertical heterogeneity of the soil is considered and the soil hydraulic parameters of the Brooks and Corey (1964) retention and hydraulic conductivity curves are calculated as continuous functions of soil texture and soil porosity using the Rawls and Brakensieck (1985) pedo-transfer function (see details in Manus et al., 2009). A

correction of soil saturated water content and hydraulic conductivity which takes into account the stones content was proposed by Manus et al. (2009). This correction is applied to both the CVN and MARINE parameters in the present study.

Figure 4a provides an illustration of the resulting variability in terms of maximum storage capacity (calculated as the sum, over all the horizons, of soil saturated water content multiplied by the horizon depth) amongst the studied catchments (Table 1). The range amongst the catchments is very large from about 60 mm to 370 mm.

The initial soil moisture is specified based on the SAFRAN-ISBA-MODCOU model with a 8x8 km² resolution as provided by Météo-France (Habets et al., 2008). For MARINE, soil parameters and initial soil moisture are distributed at the pixel scale. For the CVN model, soil parameters are distributed at the hydro-landscape scale and initial soil moisture at the catchment scale. Figures 4c and 4d show the spatial variability of initial soil water storage deficit (maximum storage capacity – initial soil water storage) for both the CVN and MARINE models. Table 1 provides a summary of the maximum and initial water storage and deficit per catchment. Initial soil water storage deficits are comparable between the models with slightly lower values for MARINE than for CVN (-7%). Figures 4c and 4d show that CVN exhibits a higher spatial variability of initial water storage deficit than MARINE due to the use of continuous pedo-transfer functions. The CVN and MARINE model are run with a zero flux bottom boundary condition.

Overland and channel flow

For the flow routing, a constant value of the Manning coefficient of 0.05 is used within the CVN model (Manus et al., 2009).

For its overland flow routing, the MARINE model requires the specification of a roughness coefficient. It is derived from a vegetation and land-use map (2000 Corine Land Cover provided by the Service de l'Observation et des Statistiques (SOeS) of the French Ministry of

Environment, www.ifen.fr). For the drainage network flow simulation within MARINE, two Manning coefficients, corresponding to the main channel and the floodplain, must be prescribed. They are taken from previous studies on the Gardons d'Anduze catchment with values of 0.05 for the Manning roughness coefficient within the main channel and 0.125 within the floodplain (Bessière, 2008).

To conclude this section, the major similarities and differences between both model structures are highlighted, based on the criteria proposed by Kampf and Burges (2007).

Spatial discretization

- The model discretizations are different. It is based on grid squares for MARINE and on homogeneous hydro-landscapes for CVN, with irregular polygonal shapes.
- Radar 1x1 km² and 5 min time step data are interpolated at the DTM grid scale within MARINE and at the hydro-landscape scale within CVN.

Surface and subsurface fluxes

- In terms of soil description, both models rely on the same soil depth and soil texture map, but the soil parameters are specified more coarsely in MARINE than in CVN (soil classes versus continuous pedo-transfer functions). Furthermore, the MARINE model assumes vertically homogeneous soils, whereas CVN accounts for vertical heterogeneity.
- Soil infiltration is described using the simplified Green and Ampt model in MARINE whereas the 1D Richards equation for saturated/unsaturated flow is used in CVN. CVN also explicitly accounts for water redistribution within the soil layers.
- Subsurface lateral flow is not considered in the intercomparison of both models

Overland and channel flow

- MARINE includes an hillslope overland flow module, based on the kinematic wave approximation of the de St-Venant equations, whereas CVN uses an empirical approach for overland flow routine
- In terms of river network, the MARINE discretization is finer and defined at the pixel scale, whereas homogeneous properties per river reach are assumed in CVN.
- The river section descriptions are also different and MARINE makes the distinction between the main channel and the floodplain, whereas only the main channel is considered in CVN. Both models use the kinematic wave approximation of the de St-Venant equations for the flow routine module.

Other processes and time discretization

- Both models are run in this study as event-based models. A variable adaptative time step is used in CVN and a fixed time step is used in MARINE
- Both models neglect interception and evapotranspiration processes during the extreme rainfall event
- Both models use the finite volume approach for the numerical solution

The two models differ in terms of spatial discretization and have quite similar approaches in terms of soil infiltration representation, with a more refined representation within CVN. Both models differ in terms of water transfer, both within the hillslopes and river network, with a more refined scheme within MARINE.

Model evaluation and sensitivity studies

The simulations are conducted from 2002/09/08 at 06hUTC to 2002/09/09 at 12hUTC, corresponding to the rainfall event duration.

Firstly, the performance of the models (reference simulations) is assessed at the regional scale using maximum peak discharge data collected during a post-flood field survey. As the models

1 are not calibrated (reference simulation), a global sensitivity analysis to model parameters is
2 conducted in order to assess the uncertainty of the model estimates. Previous global
3 sensitivity studies, using the Monte-Carlo technique, have allowed determination of the most
4 sensitive parameters in MARINE for flash floods (Le, 2008). Their number is limited using
5 the method proposed by Refsgaard (1997) where, for distributed parameters, a spatial pattern
6 is fixed and the considered parameter is a multiplicative factor of the reference map. Three
7 sensitive parameters are identified by Le (2008): a multiplicative factor for the spatial pattern
8 of soil thickness, a multiplicative factor for the spatial pattern of hydraulic conductivity and
9 the roughness coefficient of the floodplain. When lateral subsurface flow is considered, an
10 additional sensitive parameter is the multiplicative factor for the spatial pattern of saturated
11 transmissivity. Based on these conclusions, a global sensitivity analysis is performed for both
12 models for the September 2002 event using the latin hypercube method (see for instance van
13 Griensven et al., 2006 for a review of sensitivity analysis methods) with 20 intervals for each
14 parameter. Four parameters, namely a multiplicative factor for saturated hydraulic
15 conductivity, a multiplicative factor for soil depth, the Manning coefficient (main channel for
16 CVN and flood plain for MARINE, as previous sensitivity studies (Le, 2008) showed that the
17 Manning coefficient for the main channel is not sensitive in MARINE), and the initial
18 saturation are considered in the analysis. Based on expert knowledge in the region, the range
19 and distributions presented in Table 2 are assigned to the various parameters. The target
20 variables in this sensitivity analysis are the maximum specific peak discharge, time of peak,
21 and the runoff coefficient of the various catchments. The impact of rainfall description has
22 been studied in other papers (e.g. Carpenter and Georgakakos, 2004; Teztlaff and
23 Uhlenbrook, 2005; Chancibault et al., 2006; Cole and Moore, 2008; Saulnier and Le Lay,
24 2009) and is not considered here where we focus more on the impact of the description of
25 soil spatial variability and river flow.

Secondly, as the results of both models are quite similar in terms of simulation of maximum peak discharge and runoff coefficient, a comparative analysis of both model results is performed in terms of discharge and soil saturation dynamics, in order to see if differences in model behaviour can be identified. This analysis is complemented by local sensitivity studies on some parameters, and studies of the impact of various model structural choices (homogeneous versus distributed soils with CVN, inclusion of subsurface flow in MARINE).

Results

Model evaluation using post flood event data

Figure 5a shows the comparison of simulated maximum peak discharge with the field estimates (including their uncertainty). Results are summarized in Table 3 which also provides information on the peak discharge time, as derived from witness interviews when available. Figure 5a shows that MARINE simulated values are in general lower than the field observations, whereas CVN simulations tend to be higher. For both models, the peak discharge is overestimated for catchment #10 and #18. Rainfall estimation was found very uncertain for catchment #10, with krigged rainfall gauges leading to a cumulative value of 161 mm whereas the radar value was 339 mm. Catchment #18 is discussed in more detail below, as an hydrograph is available. When catchments #10 and #18 are excluded, the root mean square error is $112 \text{ m}^3 \text{ s}^{-1}$ for MARINE and $96 \text{ m}^3 \text{ s}^{-1}$ for CVN and the correlation coefficient between simulated and field values is $R^2=0.95$ for CVN with a slope of 1.14, and $R^2=0.92$ for MARINE with a slope of 0.94. The performance of both models is thus satisfactory with regards to maximum peak discharge uncertainty.

Figure 5b shows the comparison between simulated and field estimates of maximum specific peak discharge. The catchments are ordered with increasing area. We can see that maximum specific peak discharge tends to decrease with catchment area. Both models follow this trend.

MARINE (respectively CVN) estimates are within the range of uncertainty for 8 (respectively 9) catchments over 18, which is satisfactory for such events. If catchments #10 and #18 are excluded, the root mean square error is $6.4 \text{ m}^3 \text{ s}^{-1} \text{ km}^2$ for MARINE and $6.7 \text{ m}^3 \text{ s}^{-1} \text{ km}^2$ for CVN. As for maximum peak discharge, CVN tends to overestimate maximum specific discharge, whereas MARINE tends to underestimate it.

For catchment #18, a measured hydrograph is available, but the uncertainty in the measured peak is large (Table 3). The comparison of the simulated and observed hydrographs is provided in Figure 8f. The simulated hydrographs are very similar for both models. Both models overestimate the runoff volume and peak, the overestimation being larger for CVN. For this catchment, located the farthest from the radar, there may be some uncertainty on the rainfall estimate. A deeper analysis of rainfall estimates shows a cumulative rainfall of 208 mm with the radar data used in this study, 207 mm with the krigged rain gauges and only 145 mm with the data issued from the radar operational treatment. The observed runoff is 85 mm and Table 1 shows that the maximum storage capacity of the catchment is only 68 mm. Therefore, even if the catchment was completely dry, the rainfall should not exceed $85 + 68 = 153$ mm to be consistent with the observed runoff. If we take into account uncertainty in the observed runoff by applying a multiplicative factor equal to the ratio of the upper bound to the most probable peak discharge $1050/770 = 1.36$, the “observed” runoff estimation reaches 116 mm. Thus, the runoff plus storage only reaches $116 + 68 = 184$ mm, which is still below the estimated radar rainfall that we use (208 mm). The alternative hypothesis is an underestimation of catchment maximum storage capacity from the pedology map or the non-imperviousness of the bedrock.

These hypotheses and the robustness of the model estimates are examined more in details below using the global sensitivity analysis.

Global sensitivity analysis for the simulation of peak discharge and runoff coefficient

The results of the global sensitivity analysis in terms of maximum specific discharge and runoff coefficient are summarized in Figure 6 for both models. Almost none of the parameter set is able to simulate the observed values of maximum specific discharge for catchments #10 and #18. It gives more credence to the hypothesis of a problem with rainfall estimation. For catchment #18, only two parameters sets, associated with multiplicative factors of 8.1 (resp. 2.1) for saturated hydraulic conductivity, and of 3.6 (resp. 9.4) for soil depth, associated with dry initial conditions (0.11 and 0.34 respectively) are able to provide Nash efficiency values larger than 0.7 for both models. The high values of these multiplicative factors also questioned the relevance of the information provided by the soil data base about soil depth and the pedo-transfer functions used for the estimation of soil hydraulic properties. Enhanced *in situ* measurements of these properties are therefore required.

Figure 6 confirms the tendency observed in Figure 5 of an overestimation of maximum peak discharge by CVN and an underestimation by MARINE. For both models, the uncertainty range of the maximum specific peak discharge is in general within the error bounds of the post field data. For the catchments which are poorly simulated by CVN (#3, 7, 8, 10, 15 to 18) or MARINE (#1, 4, 5, 8, 10, 12, 16 to 18), the parameter uncertainty bounds are generally outside the errors bounds of the observed values. In terms of peak discharge (not shown), the time amplitude is small (less than 20 min, except when there are several peaks). When there is a peak of high rainfall intensities (larger than 100 mm hr^{-1}), the time of the peak is very stable, showing that these high rainfall intensities are probably very influential on the peak discharge, especially for the CVN model. For the CVN model, spurious rainfall intensity peaks might be the reason for the general overestimation on the badly simulated catchments. These rainfall peaks are likely to generate infiltration excess runoff, which is instantaneously transferred to the river network. Within the MARINE model, overland flow can smooth the impact of these

1 rainfall peaks. For CVN, the reference simulation is in general very close to the median
2 (Figure 6a), whereas it is closer to the 25th quartile for MARINE.

3 In terms of runoff coefficient, the similarity between both model response is large (Figure 6c
4 and d). The large values of the median runoff coefficient for both models must be highlighted.

5 The CVN 95% interval for the runoff coefficient is in general closer to a runoff coefficient of
6 1 than MARINE. Only some simulations (in general associated with the largest multiplicative
7 factors for soil depth) lead to small values of the runoff coefficient. Figure 7b to d show the
8 relationship between the runoff coefficient and the multiplicative factors for saturated
9 hydraulic conductivity, soil depth and with initial saturation for catchment #15, which is
10 representative of the other catchments. Both models present similar patterns, although the
11 absolute values are different (see next section). This shows that, despite the differences in
12 terms of surface and subsurface representation, the signature of the difference in model
13 structure is weak in terms of runoff coefficient patterns. For the September 2002 extreme
14 event, with an exceptionally large rainfall amount, a generalized soil saturation is simulated
15 for most of the parameter sets. Neither the maximum peak discharge, nor the runoff
16 coefficient allow the identification of significant differences in responses corresponding to
17 differences in model structures for the runoff production components.

18 Figure 7a illustrates the dependence of the simulated maximum peak discharge as a function
19 of the Manning coefficient for catchment #15, but the pattern is similar for the other
20 catchments. For both models, when the Manning coefficient is increased (in log scale), an
21 linear decrease of the maximum peak discharge is simulated. The slope of the relationship is
22 smaller for MARINE than for CVN, showing a lower sensitivity of MARINE to this
23 parameter. This also reflects the fact that MARINE distinguish between the main channel and
24 the floodplain. The reference simulation values are quite different between both models,
25 reflecting different choices in the flow routing and parameter values. When compared to the

observed maximum peak discharge (horizontal lines), a range of more plausible values can be highlighted. But the results are different according to the catchments (not shown). The determination of a regional value for the Manning coefficient, based only on the maximum peak discharge appears difficult.

In conclusion, given the available information and the uncertainty, both models perform satisfactorily for a wide range of catchment sizes and rainfall input amounts. The results of the global sensitivity analysis show that differences in model structures are not reflected in differences in runoff coefficients. They are partly reflected in differences in maximum specific discharge, but this variable is not sufficient to assess the most suitable model structure. In the remaining of the paper, we go a step further by comparing the model results in terms of simulated process dynamics (discharge and soil saturation) in order to see if information on other variables could help determining which model structure better represents actual processes. The reference simulation is used for this purpose.

Nevertheless, we would like to stress the value of post flood field data of maximum peak discharge, as they are the only data available to assess model performance at the regional scale and for catchment areas less than 100 km². In Bonnifait et al. (2009), their Figure 6 shows that in the studied region, gauged catchments are only available for areas larger than 100 km². Observed hydrographs are of course valuable to verify the simulated flow dynamics for extreme events. However, discharge estimation remains very uncertain for extreme events as stage-discharge relationships are often extrapolated far beyond the maximum gauged value ,and only a few locations are gauged.

Catchment hydrological response and flow dynamics

Figure 8 provides the simulated hydrographs by both models for some selected catchments. The horizontal bars show the post flood field estimates of peak discharge and their

1 uncertainty. When available, the times of the peak as stated by witnesses are shown as
2 crosses. Table 3 shows the comparison between simulated and field estimates of peak
3 discharge and time of peak for both models. Figure 8 shows that the shape of the simulated
4 discharge is highly related to the rainfall dynamics with peaks in general associated with high
5 rainfall intensities. The CVN model tends to be more reactive to rainfall than the MARINE
6 model and shows more pronounced peaks than MARINE, with sometimes a large
7 overestimation (#3, #7, #8, #16). It should be noted that MARINE also overestimates the
8 peaks for these catchments (except #3). Differences in model responses increase with
9 increasing size of the catchment (Figure 8) with less pronounced peaks with MARINE than
10 with CVN as long as the catchment area increases. The differences in the transfer modules
11 lead to a larger smoothing of the hydrological response in MARINE than in CVN when the
12 catchment size increases. These differences can be partly related to the transfer module and
13 the difference in the Manning coefficient as shown in Figure 9 and Table 4, and confirmed by
14 the global sensitivity analysis (Figure 7a). When the Manning coefficient is increased from
15 0.05 to $0.125 \text{ m}^{-1/3} \text{ s}^{-1}$, the peak discharge decrease by 16-36% in CVN and by 7-30% in
16 MARINE. The time of peak is delayed by 10-30 min in CVN and by 20-30 min in MARINE.
17 The case of #17 is particular: two peaks are simulated and the change in roughness produces a
18 change of the maximum peak in MARINE which leads to an 11h30 hour difference. Figure 9
19 and Table 4 show that, in general, peaks occur in MARINE before they do in CVN when
20 both models use the same roughness. This is certainly a consequence of the differences in
21 flow routing representations (as both models behave similarly in terms of runoff generation).
22 Figure 9 also shows that with identical roughness coefficients, the response of MARINE
23 remains smoother than that of CVN. In particular, the decrease in peak discharge is less
24 pronounced in MARINE which may be a consequence of the accounting of hillslope transfer
25 within MARINE.

When information about the peak hour is available, model estimates agree fairly well with the values provided by the witnesses (or observation for Saumane), with a difference of less than 30 min for #1 (second peak), #6, #9, #17 and #18. Differences in peak time estimates can reach more than one hour for #1 (first peak), #3 and #11. The dynamics of #6 and #9 are particularly well represented by both models (Figure 6c). But these catchments have a very shallow soil (Table 1) and are almost completely saturated on 09/08 at 12UTC, after only 6 hours of rainfall (see next section). Thus, the observed peaks are mainly explained by the rainfall intensity peaks leading to saturation excess runoff. Both models well represent this dynamic, showing that, for these small catchments (about 10 km²), water transfer was satisfactorily represented.

As also shown by the global sensitivity analysis, for the different catchments, the simulated peak hours by both models are quite similar although the transfer modules are different. However, the amplitude of the peaks is much higher in CVN than in MARINE. Generally, the sensitivity of MARINE to the Manning roughness coefficient is smaller than that of CVN. This is likely because of channel flow modelling which accounts for drainage network within floodplains. More detailed information about the river bed, such as its geometry and roughness would help improve the transfer modules of both models. The availability of distributed water height series (and preferably discharge series) at various scales could also be helpful in better characterizing flow dynamics.

Table 3 also provides the runoff coefficient, RC, simulated by both models. The average difference between MARINE and CVN RC estimation is 0.04 in the range [-0.01, 0.15] and the root mean square difference is 0.07. The difference between both models in terms of runoff coefficient is thus much higher than expected given the relative proximity in their process representation. These differences in hydrological response also appear in Figure 10

which shows the relationships between the simulated RC and rainfall total amount, average soil depth, initial storage deficit and initial saturation. On this graph, we have distinguished between points with an acceptable or a bad simulation of maximum peak discharge. Figure 10 shows no significant correlation with all the variables for MARINE ($R^2 < 0.20$), whereas a significant relationship is found between CVN RC and average soil depth ($R^2 = 0.62$, $p < 10^{-6}$, Figure 10b) and initial soil water deficit ($R^2 = 0.75$, $p < 10^{-6}$, Figure 10c). These relationships remain valid when only the “good” simulations are taken into account. In CVN, the strong relationship between runoff generation, soil depth and initial storage deficit is expected as the generated ponding is highly related to soil depth and/or water storage deficit in the case of full saturation. As the ponding is directly routed to the river, all the ponding volume finally reaches the outlet after being transferred within the river. The larger scatter for MARINE could be explained by the hillslope routing which allows some re-infiltration before the water reaches the river.

Soil saturation dynamics

Figure 11 maps the saturation state of each catchment at several times (from top to bottom: 8 September 2002 at 12hUTC, 9 September 2002 at 00hUTC, 9 September 2002 at 12hUTC). The spatial distribution of the saturation states is quite similar for both models; however CVN simulation results in higher saturation, as can be seen in Figure 12. The final saturation is 0.96 for CVN while it is 0.90 for MARINE. This difference may be related to CVN’s description of the vertical soil heterogeneity as explained in the model set up section. MARINE exhibits lower values of initial soil water storage deficit for deeper soils (Figure 4). However, the saturated hydraulic conductivity used in MARINE is generally lower than the one used by CVN with a ratio between both models ranging between 0.25 and about 9 (not shown). This may explain both the lower final saturation and the higher RC obtained with MARINE. Table 2 shows that MARINE runoff coefficients are lower than CVN RC only for 2 catchments.

To illustrate the role of soil properties on the soil saturation dynamics, the mean saturation state is plotted against different soil characteristics: soil classes (Figure 13a) and soil depths (Figure 13b). Results show that soil spatial variability has a great impact on this dynamic. Indeed, for both models, soils presenting high hydraulic conductivities (13.2 mm h^{-1} for class n°4) or low depths (ranging between 0 m and 0.2 m) are rapidly saturated. The important increase of saturation between September 8th 2002 at 18hUTC and September 9 at 00hUTC can be explained by the accumulated rainfall during these 6h : locally more than 200 mm on catchments presenting class 4 soil texture (#12, #6), and locally more than 150 mm on catchments presenting low soil depths (#9, #11). For these soils, CVN and MARINE exhibit almost the same saturation dynamics: the vertical soil heterogeneity is indeed of little importance in these cases. CVN presents higher saturation states than MARINE for soils with low hydraulic conductivities (0.6 mm h^{-1} for class n°11) or large depths (ranging between 0.8 m and 1 m) where soil can present high vertical heterogeneity. The important difference found for soil class n°11 may also be explained by the soil parameter specifications of each model: soil class n°11 covers an important range of soil textures (with a clay content ranging from 40 to 60%) with different values of soil parameters in CVN as continuous pedo-transfer functions are used, while soil parameters are fixed for each soil class in MARINE. Indeed, it can be seen in Figure 14b that the time evolution of the mean saturation states is highly related to the soil class and therefore to hydraulic conductivity in MARINE, whereas it is more related to soil depth (Figure 14a) in CVN. All these results emphasize the impact of the spatial distribution of rainfall and soil properties on simulated soil saturation dynamics.

Figure 15 shows typical infiltration dynamics as simulated by the CVN models for different soil STUs. Figure 15 shows that for the shallowest soils, full saturation is reached on 09/08 at 12hUTC on US 430 (Figure 15a) and US 425 (Figure 15c), with the generation of infiltration excess runoff related to a low saturated hydraulic conductivity ($K_s=0.7 \text{ mm hr}^{-1}$). Even if

shallow, US 535 is saturated later on 09/09 at 00hUTC (Figure 15b) due to a higher $K_s=14$ mm hr⁻¹. These three soils are fully saturated on 09/09 at 00hUTC, leading to saturation excess runoff until the end of the event. With the no flux boundary condition, these profiles do not change until the end of the simulation within the CVN model. For US 528, a moderate value of $K_s=4$ mm hr⁻¹, and a deep soil lead to a topsoil saturation, and the generation of infiltration excess runoff, on 09/09 at 00hUTC (Figure 15d), but infiltration can continue into deeper layers until 09/09 at 12hUTC when the soil is fully saturated, generating saturation excess runoff. Finally US 529 has a low $K_s=0.7$ mm hr⁻¹ and two horizons. The first horizon is saturated on 09/09 at 18hUTC, but infiltration continues into the second horizon (Figure 15e). Infiltration excess runoff is generated throughout the whole event, but at the end, the soil is not fully saturated and saturation excess runoff has not occurred in this soil. If US 529 is assumed vertically homogeneous (Figure 15f), the topsoil saturation occurs later than with the two horizons on 09/09 06hUTC, and the soil is not fully saturated at the end of the event.

Impact of soil description on the simulated hydrographs

Sensitivity tests exploring the impact of soil spatial variability on the CVN simulated response are presented in this section for three catchments: #15, #17 and #18. Catchment #15 and #17 were chosen for this sensitivity study because they have a large variability of soils with different storage capacity (Figure 4a) and hydraulic conductivities.

Simulations are performed using the same initial saturation rate as those given in Table 1. Additional simulations are performed by using the same initial water storage deficit. Results are shown in Figure 16a, b and c and Table 5. The impact on peak discharge is moderate ($\pm 7\%$ for #15 and #17 and $+13\%$ for #18), but the impact on runoff is significant with a variation of ± 0.10 in the runoff coefficient. In general the changes in runoff are related to the changes in initial soil water deficit and the increase/decrease in runoff amounts is equal to the

increase/decrease in initial storage deficit. When the same initial storage deficit is used for initial conditions, runoffs are very similar, but there is an impact on peak discharge of $\pm 7\%$. When soils are assumed to be vertically homogeneous instead of taking into account the various horizons, the impact is limited to variations in the runoff coefficient and peak discharge of a few percent. As does the previous analysis of soil saturation dynamics, this sensitivity study highlights the importance of the initial water storage deficit on the hydrological response at the outlet. Note that the quantification of initial conditions using a saturation degree can lead to large differences in initial water storage deficit and is thus not recommended. However, quantifying the initial soil water deficit is much more demanding in terms of landscape description as it requires an accurate documentation of soil depths and porosity.

Impact of including subsurface flow within MARINE

In order to test the impact of subsurface lateral flow within the MARINE model, a simulation is performed with the same parameters as the reference simulation and activation of subsurface lateral flow at a uniform subsurface velocity of 2 m.h^{-1} . This value is chosen as 50 times the averaged Green and Ampt hydraulic conductivity over the whole studied area. Results show little sensitivity of peak discharge and runoff coefficient to subsurface flow with a maximum difference of respectively -1% and $+0.03$ for #18 catchment. No impact on peak time is observed. It seems that for such an extremely intense and short event, slow subsurface lateral flow is of little importance in simulating peak time and discharge: with a simulation duration of 30h and a subsurface velocity of 2 m.h^{-1} , no significant contribution of subsurface flow to outlet hydrographs is simulated. This conclusion should however not be generalized to less intense events. Indeed, introduction of subsurface flow slightly modifies the recession simulation ($6 \text{ m}^3 \text{ s}^{-1}$ difference, hardly visible in Figure 16d), which may be of importance for longer events with several rainfall peaks for instance. Moreover, the inclusion of subsurface

lateral flow has an important impact on distributed saturation dynamics. Indeed, Figure 17 maps the saturation state of catchment #18 at the end of the simulation (September 9th 2002 at 12hUTC) with and without subsurface flow. It clearly shows the importance of the drainage network in the evolution of the saturation state when subsurface flow is considered. This is due to the exfiltration that can occur in the drainage network. Measures describing the spatial distributions of saturation state would be valuable to chose which processes must be accounted for in the model.

Discussion

In this study, two distributed hydrological models are set up at the regional scale using the available information and expert knowledge, without specific calibration. They are run for the September 8-9 2002 event in the Gard on 18 sub-catchments ranging from 2.5 to 99 km² where data of maximum peak discharge, and sometimes time of peak are available. Post flood estimates of maximum peak discharge and time of peak are used to evaluate the models at the regional scale. This kind of information is very valuable because it is the only information available for small catchments. Up to now few studies (e.g. Bonnifait et al., 2009; Saulnier and Le Lay, 2009) have targeted the regional scale in the study area. These studies were based on a calibrated Topmodel approach. Bonnifait et al. (2009) also used the post flood field data for model evaluation but the comparison was restricted to catchments larger than 50 km². The originality of our study is to tackle very small catchments, which were shown to be the most vulnerable to flash floods (Ruin et al., 2008).

The comparison between field estimates and simulated peak discharge shows a reasonable agreement for half of the catchments for both models. The global sensitivity analysis on three soil parameters and the Manning coefficient shows that the results of the reference simulation in terms of maximum specific peak discharge are close to the median of the parameter set for CVN and to the 25th quartile for MARINE. The uncertainty range of CVN are generally larger

1 than that of MARINE. Further analysis shows that CVN is more sensitive to the Manning
2 roughness coefficient of the river bed than MARINE (for which it is the floodplain roughness
3 coefficient) with a decrease of peak discharge when the Manning roughness coefficient is
4 increased. Peak discharge could be used to reduce the range of plausible Manning roughness
5 coefficients but, for the studied event, it seems difficult to find a unique value suitable for all
6 the catchments. For some catchments, the sensitivity study highlights probable problems with
7 the rainfall amount, as almost none of the parameter set is able to simulate the field peak
8 discharge. The time of peak is found quite stable with regards to the parameters and most of
9 the time, it is related to the time of high rainfall intensity peaks (larger than 100 mm hr^{-1}). The
10 CVN model is more sensitive to these peak intensities than the MARINE model. In CVN such
11 peak rainfall intensities generate infiltration excess runoff which is directly transferred to the
12 closest river reach, leading to large peak discharge. In MARINE, overland flow allows
13 possible re-infiltration of this infiltration excess runoff, and a smoothing of the impact of
14 those peaks. The corresponding peak discharge are therefore lower. Post flood estimates of
15 maximum discharge are therefore useful for analysing the impact of various flow routine
16 modules. However, they do not allow the evaluation of the simulated flow dynamics and the
17 identification of the most relevant model response. For this purpose, regional values of water
18 height (and possibly peak discharge) are required in order to base the analysis of the whole
19 hydrograph response. Our analysis should also be complemented by a study of the impact of
20 river geometry on the results.

21
22 In terms of runoff coefficient, post flood maximum discharge estimate do not provide a mean
23 to assess the relevance of the simulated results nor to discriminate between the two model
24 structures. The global and local sensitivity analysis highlight the large impact of soil depth,
25 saturated hydraulic conductivity and initial soil moisture on the simulated response, especially

1 in terms of runoff coefficient. The global sensitivity analysis also shows that several
2 combinations of these parameters can lead to similar responses in terms of peak discharge or
3 runoff coefficient. A better knowledge – at the regional scale- of soil depths, hydraulic
4 properties and initial conditions is therefore required in order to reduce uncertainty in these
5 parameters and in model results (Loague and VanderKwaak, 2004; Ebel and Loague, 2006).
6 For the extreme September 2002 event, very high runoff coefficient are simulated by both
7 models. Differences in model responses, related to the different model representation can be
8 highlighted but they are masked by the importance of the event.

9
10 In the present study, evapotranspiration and interception processes are neglected. This
11 hypothesis is reasonable given the duration of the event (one to two days) and the high rainfall
12 intensities during the event. Vegetation and land use is considered within MARINE for the
13 specification of the Manning roughness coefficient for overland flow, but a previous
14 sensitivity analysis showed that this parameter is not very influential (Le, 2008). On the other
15 hand, vegetation and land use do have a significant impact on initial conditions. A more
16 comprehensive version of the CVN model is under construction. It will be used to study the
17 impact of soil moisture. A continuous model is considered and an evapotranspiration module
18 is being added to the current version of the model with parameters depending on the land use
19 and vegetation.

20
21 Several studies (e.g., Carpenter and Georgakakos, 2004; Teztlaff and Uhlenbrook, 2005;
22 Chancibault et al., 2006; Cole and Moore, 2008; Saulnier and Le Lay, 2009) highlight the
23 importance of an accurate quantification of the rainfall spatio-temporal dynamics for an
24 accurate simulation of the catchment's response. This point is not explicitly addressed in our
25 study. However, some of our results about time of peak show the importance of high rainfall

1 intensity peaks on the modelled results. The highest sensitivity of the CVN model to those
2 rainfall peaks is certainly related to its very simple overland flow representation. This
3 component of the model can be improved in the future. For some catchments where problems
4 are encountered with both models, rainfall accuracy is questioned. Therefore, the study
5 confirms the need for further work on the improvement of high spatial and temporal rainfall
6 resolution, and on the appropriate location of rainfall with regards to catchments boundaries.
7 Although the September 2002 event is perhaps too extreme to fully illustrate this point – due
8 to a generalized soil saturation- our study also highlights the importance of the representation
9 of soil spatial variability, both in terms of soil depths, porosity and hydraulic properties.
10 Although infiltration excess runoff is generated for all the catchments due to the high rainfall
11 intensities, the soil water storage deficit is found to be very important when it comes to the
12 timing of the response. This is in particular related to the threshold effect of full saturation of
13 the soils. As a consequence, information about the initial soil water deficit is more relevant
14 for the characterization of the initial moisture conditions than is the saturation degree of the
15 catchments, as provided by the SAFRAN-ISBA-MODCOU model. However, the former
16 estimation requires more data than the latter, as soil depths and porosity must be known with
17 good accuracy. When subsurface lateral flow is considered, its impact on such an extreme
18 event is found to be negligible in terms of peak discharge and runoff volume. However this
19 conclusion must not be generalised to less intense events.
20 The saturation dynamics analysed at the regional scale, provide interesting insight into
21 hydrological processes. When subsurface flow is not activated, the soil saturation patterns are
22 mostly related to soil depth (and rainfall), whereas, when subsurface flow is considered, a
23 second pattern of organisation along the river network is observed. Such patterns are typical
24 of models based on the Topmodel concepts (e.g. Quinn et al., 1998). Patterns of soil moisture
25 have been reported for small catchments using in situ data (e.g. Western et al., 2004; Latron

and Gallart, 2007), but a much larger coverage would be required to assess the relevance of our simulated patterns and the methodology proposed by Beldring et al. (1999) could be used. Remote sensing of soil moisture has proven useful in reducing model parameter uncertainty in hydrological models (e.g. Franks et al., 1998) and has been shown to improve their performance (e.g. Pauwels et al., 2001). However, remote sensing only provides retrieval of topsoil moisture whereas an integrated value over the whole profiles would be required. Works are conducted to retrieve integrated soil moisture over the root zone from the combination of surface soil moisture and soil vegetation atmosphere transfer modelling (e.g. Calvet and Noilhan, 2000; Montaldo et al., 2003; Albergel et al., 2008). The application of such techniques on the studied region should be evaluated, because information about integrated soil moisture content is required to assess which of our simulated patterns is the most realistic.

The general methodology proposed in this paper is the following i) set up of distributed hydrological models at the regional scale using the available data and expert knowledge without specific calibration; ii) a regional evaluation using the available information such as available hydrographs, post flood data of maximum peak discharge and time of peak; iii) sensitivity studies to input data, parameters and process representation in order to get insight into the most sensitive points and derive information about processes, data and parameters requiring further measurements and understanding. It is applied to the September 2002 extreme event where the required data is available. Some interesting conclusions, summarized below are drawn from this study. However, the event is so extreme that a general saturation of the catchment is simulated, and the impact of various modelling hypotheses cannot be fully examined. The study should be repeated with less extreme events (Merz and Plate, 1997) in order to get more contrasted responses between catchments and models. The methodology can also be applied to other regions, provided the required data are available.

Conclusions

In the introduction, four questions are highlighted. The results presented in this paper provide some interesting responses. The study shows that it is possible to set up , at the regional scale, distributed hydrological models, relevant for flash flood simulation, using available data and expert knowledge. It is a step towards the “model of everywhere”, advocated by Beven (2003). Provided high spatial and temporal resolution of rainfall is available, post flood field data of maximum peak discharges and time of peak are valuable for the regional evaluation of these models. However, they are not sufficient for a full assessment of the simulated hydrological response in terms of flow dynamics, runoff coefficient and soil saturation dynamics. Model sensitivity studies including input data, parameters and process representation are very valuable in order to highlight where additional data could be useful.

This point is further detailed below where an experimental set up, suitable for improved understanding and modelling of flash floods is proposed in the context of the future HyMeX program. This experimental set up is based on nested instrumented catchments at various scales: i) densely instrumented small catchments of 1-10 km², where process studies can be conducted; ii) larger instrumented catchments (of about 100-500 km²) where the change of scale problem can be tackled and, iii) regional catchments (of about 1000-10000 km²), where operational data are collected. Such a strategy is already implemented in African within the AMMA project (Lebel et al., 2009), or in the UK in the framework of the Catchment Hydrology And Sustainable Management program (O’Connell et al., 2007). For flash floods studies, the following recommendations can be made:

- The information provided by post event field experiments about maximum peak discharge and time of peak is definitively recognised as very valuable, especially for small

catchments. But the availability of enhanced networks of discharge series and/or water levels would increase our capability to specify hydrological models parameters and improve the process representations, through a better knowledge of the flow dynamics. The set up of LS-PIV (Large Scale Particle Image Velocimetry) networks (Muste et al., 2005; Le Coz et al., 2009²) and/or water level measurement networks (Sarrazin et al., 2009) which increase the density of this information in nested catchments is therefore encouraged.

- The spatial variability of soil depth and soil porosity, as well as hydraulic properties is found to be influential on the hydrological response and more specifically on the soil saturation patterns and dynamics. Up to now, this variability has been described through pedological maps, which are built for agronomic purposes, but not necessarily relevant to hydrological purposes since we are interested in the whole soil profiles, including the potentially altered substratum. The description of soil properties such as soil depth, porosity and hydraulic conductivity should thus be enhanced at the regional scale.
- Finally this study proposes different patterns of soil saturation dynamics according to different model hypotheses about the role of subsurface lateral flow. Regional mapping of soil saturation, over the whole profiles, would be required to determine which patterns are the most realistic as compared to field conditions. A combination of remote sensing data, in situ local measurements, and data assimilation could provide the necessary data.

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² Le Coz, J., Hauet, A., Dramais, G., Pierrefeu, G., 2010. Improvement of flash-flood discharge measurements using image-based velocimetry (LS-PIV), Journal of Hydrology, Flash Floods special issue, accepted.

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List of figures

Figure 1. Location of the study area, location of the simulated catchments and slope map of the region. The catchment numbers are those of Table 1. Coordinates are in extended Lambert II (m)

Figure 2. Scheme of the MARINE model. a) MARINE model structure. b) Geometry of a network cross section within the MARINE model

Figure 3. Scheme of the CVN model. Step 1 derives the hydrographic network and sub-catchments boundary from the DTM. The river network provides the geometry for the 1D flow routine module. Step 2 derives the hydro-landscapes, used as elementary modeling unit for the Richards equation infiltration module, from the overlay of the pedology map and sub-catchments boundaries. The hydro-landscapes are sub-divided into layers to account for soil vertical heterogeneity and ponding is directly transferred to the closest river reach.

Figure 4. a) Maximum soil water storage capacity (mm) for the studied catchments. b) Total cumulated rainfall (mm) between 2002/09/08 06hUTC and 2002/09/09 12hUTC. c) Initial soil water deficit (mm) for the CVN model on 2002/09/08 06hUTC. d) Initial soil water deficit (mm) for the MARINE model on 2002/09/08 06hUTC.

Figure 5. a) Comparison of simulated (● for CVN and ▲ for MARINE) and field estimation of maximum peak discharge in a log-scale. The vertical bars along the 1:1 lines correspond to the uncertainty in the estimated values. b) Comparison of simulated (● for CVN and ▲ for MARINE) and field estimation (□ with the corresponding error bar) maximum specific peak discharge. The x-axis provides the catchment numbers (see Table 1) and the catchments are ordered with increasing area.

Figure 6. Boxplot of the global sensitivity analysis results on maximum specific peak discharge (top) and runoff coefficient (bottom) for the CVN model (left) and MARINE model (right). The thick line in the boxplot represents the median, the limits of the box provide the

25th and 75th quartiles and the whiskers provide the 95% confidence interval. Crosses are the outliers. For maximum specific discharge, the black points with the thick vertical lines are the *in situ* field estimates with their uncertainty range (see Table 3)), and the black squares are the results of the reference simulation.

Figure 7. Most significant results of the sensitivity analysis for catchment #15 (grabieux). (a) Maximum specific discharge as a function of Manning roughness coefficient ($m^{-1/3} s$). The horizontal full line is the *in situ* field estimate with its uncertainty (dashed line). Runoff coefficient as function of the (b) multiplicative factor for saturated hydraulic conductivity; (c) soil depth; and (d) of initial soil moisture. CVN results appear as black squares and MARINE results as open triangles. The larger symbols correspond to the reference simulation.

Figure 8. Examples of simulated hydrographs by the MARINE and CVN models for catchments (a) #1, (b) #3, (c) #9, (d) #15, (e) #17, (f) #18. The horizontal full lines are the field estimated peak discharge and the horizontal dashed lines are the confidence intervals. Crosses are drawn at the time of peaks as retrieved from witness interviews. The observed hydrograph for catchment #18 is shown with black squares on panel (f). Catchments are presented in the figure in increasing area order.

Figure 9. Sensitivity study to the specification of the Manning coefficient of the CVN model (reference value: $n=0.05 m^{-1/3} s$, tested value: $n=0.125 m^{-1/3} s$) and of the floodplain Manning coefficient of the MARINE model (reference value: $n=0.125 m^{-1/3} s$; tested value: $n=0.05 m^{-1/3} s$) for (a) #10; (b) #15; (c) #17; (d) #18.

Figure 10. Relationship between catchment simulated runoff coefficients and (a) total rainfall amount ; (b) average soil depth ; (c) initial soil water deficit (mm) ; (d) initial saturation ratio. The graphs distinguished catchments for which the simulation of maximum peak discharge is within the error bounds (“good”) and out the error bounds (“bad”).

Figure 11. Maps of the simulated soil saturation state by CVN (left) and MARINE (right) for three dates: 09/08 12hUTC (top), 09/09 00hUTC (middle) and 09/09 12hUTC (bottom).

Figure 12. Time evolution of the average saturation state over the whole studied area for both models.

Figure 13. Example of time evolution of the mean saturation state for different types of soils a) Rawls and Brakensiek 4 and 11 soil classes (4: loam; 11: clay); Soil depths ranging between 0-0.2 m and 0.8-1 m are distinguished.

Figure 14. Correlation of the saturation dynamics with soil properties at several times (square: 8 september at 12:00; triangle: 9 september at 00h; diamond: 9 september at 12h) a) Soil depths for CVN; b) Soil classes for MARINE (4: loam; 7: clay loam; 11: clay - Only soil classes present in more than 1000 cells of the grid have been taken into account).

Figure 15. Time evolution of the soil saturation profiles of six different soils on different catchments. (a) #15 US 430; (b) #15 US 535; (c) #17 US 425; (d) #15 US 528; (e) #17 US 529; (f) #17 US 529 vertically homogeneous. The symbols correspond to the various dates: (full line) 2002/09/08 06h00; (■) 2002/09/08 12h00; (▲) 2002/09/08 18h00; (□) 2002/09/09 00h00; (■) 2002/09/09 12h00; (+) 2002/09/09 12h00.

Figure 16. (a) Sensitivity to the soil representation on #15 for the CVN model; (b) Sensitivity to the soil representation on #17 for the CVN model (c) Sensitivity to the soil representation on #18 for the CVN model; (d) Sensitivity of the MARINE model to inclusion of sub-surface flow on #18

Figure 17. Spatial distribution of the saturation state of #18 catchment on September 9 2002 at 12hUTC simulated with MARINE model a) without subsurface lateral flow b) with subsurface lateral flow at a uniform velocity of 2 m.h^{-1} .

Table 1. Characteristics summary of the studied catchments, ordered by increasing area.

#	Catchment name	Area (km ²)	Average soil depth (cm)	CVN maximum storage capacity (mm)	MARINE maximum storage capacity (mm)	Initial saturation (%)	CVN initial storage deficit (mm)	MARINE initial storage deficit (mm)
1	alzon_sec05	2.5	49	155	146	54	71	67
2	alzon_sec04	3.4	21	94	88	54	43	40
3	droude_sec04	4.0	84	374	352	55	169	158
4	braune_sec06	7.3	45	161	148	58	67	61
5	alzon_sec01	8.2	25	109	98	54	50	45
6	ourne_sec03	10.2	25	104	98	58	43	41
7	braune_sec05	11.6	43	145	137	57	62	59
8	bourdic_sec02	12.0	45	145	97	62	43	41
9	ourne_sec02	12.0	26	103	137	58	55	52
10	braune_sec01	14.6	73	223	208	58	92	86
11	alzon_sec03	16.0	30	123	113	54	56	52
12	crieulon_sec01	19.0	96	397	366	68	128	117
13	galeizon_sec05	21.0	18	62	59	48	33	31
14	braune_sec04	23.3	64	249	231	59	102	95
15	grabieux_sec02	24.1	45	185	168	56	80	73
16	galeizon_sec03	38.1	22	72	68	47	37	36
17	courme_sec04	50.2	81	336	314	61	131	123
18	saumane	99.0	19	68	53	46	36	29

Table 2: Distributions and range of the parameters used in the global sensitivity analysis

Parameter	Range	Type of distribution	Comment
Multiplicative factor for saturated hydraulic conductivity	[0.01, 100]	log-normal	Local in situ field measurements report high saturated hydraulic conductivity (Marchandise, 2007)
Multiplicative factor for soil depth	[0.01, 10]	Log-normal	
Manning coefficient (main channel for CVN, floodplain for MARINE)	[0.01, 0.2]	Log-normal	From the literature
Initial soil saturation	[0,1]	uniform	

Table 3. Estimated and simulated maximum peak discharge. Estimated and simulated time of peak. When two values are provided, it corresponds to two successive peaks. Simulated runoff coefficients (RC) for the MARINE and CVN models.

#	Estimated peak discharge ($\text{m}^3 \text{s}^{-1}$)	Estimated time of peak (UT).	Total accumulated rainfall (mm)	CVN simulated peak discharge ($\text{m}^3 \text{s}^{-1}$)	MARINE simulated peak discharge ($\text{m}^3 \text{s}^{-1}$)	CVN simulated time of peak	MARINE simulated time of peak	CVN runoff coefficient	MARINE runoff coefficient
1	100 [80 - 125]	09/09 01:00-02:00 09/09 05:00-06:00	526	109	68	09/09 00:10 09/09 05:40	09/09 00:15 09/09 05:10	0.90	0.92
2	100 [70 - 120]	-	393	85	72	-	-	0.88	0.90
3	40 [30 - 50]	08/09 22:00 09/09 05:00-06:00	368	95	56	08/09 21:00 09/09 08:10	08/09 20:40 09/09 07:20	0.72	0.87
4	160 [120 - 200]	-	441	185	91	-	-	0.84	0.85
5	330 [270 - 370]	-	436	269	207	-	-	0.87	0.89
6	270 [220 - 350]	09/09 02:00 09/09 05:00	502	272	226	09/09 02:10 09/09 05:40	09/09 02:00 09/09 05:40	0.91	0.92
7	230 [170 - 290]	-	420	390	170	-	-	0.85	0.92
8	111 [100 - 111]	-	371	213	156	-	-	0.85	0.94
9	300 [250 - 350]	09/09 02:00 09/09 05:00	505	333	271	09/09 02:10 09/09 05:40	09/09 02:00 09/09 05:50	0.91	0.91
10	60 [40 -]	-	339	419	163	-	-	0.79	0.87
11	430 [300 - 550]	09/09 03:00-03:30	480	532	376	09/09 05:40	09/09 05:40	0.89	0.90
12	320 [285 - 380]	-	451	357	217	-	-	0.78	0.85
13	400 [320 - 490]	-	336	452	375	-	-	0.90	0.89
14	300 [200 - 400]	-	429	359	320	-	-	0.77	0.83
15	390 [310 - 470]	-	404	522	348	-	-	0.79	0.82
16	400 [350 - 500]	-	267	645	616	-	-	0.86	0.88
17	635 [590 - 730]	09/09 10:00	360	539	307	09/09 09:30	09/09 9:40	0.75	0.89
18	770 [650 - 1050]	09/09 05:00	207	1673	1175	09/09 04:30	09/09 04:20	0.84	0.83

Table 4. Sensitivity study to the specification of the Manning coefficient of the CVN model (reference value: $n=0.05 \text{ m}^{-1/3} \cdot \text{s}$, tested value: $n=0.125 \text{ m}^{-1/3} \cdot \text{s}$;) and of the floodplain Manning coefficient of the MARINE model (reference value: $n=0.125 \text{ m}^{-1/3} \cdot \text{s}$; tested value: $n=0.05 \text{ m}^{-1/3} \cdot \text{s}$;) for 4 catchments.

#	CVN peak discharge ($\text{m}^3 \text{ s}^{-1}$)	CVN time of peak (UTC)	CVN peak discharge ($\text{m}^3 \text{ s}^{-1}$)	CVN time of peak (UTC)	MARINE peak discharge ($\text{m}^3 \text{ s}^{-1}$)	MARINE time of peak (UTC)	MARINE peak discharge ($\text{m}^3 \text{ s}^{-1}$)	MARINE time of peak (UTC)
<i>n</i>	0.05		0.125		0.05		0.125	
3	95	09/08 21h00	65	09/08 21h10	71	09/08 20h40	65	09/08 21h10
15	516	09/09 05h50	375	09/09 05h50	435	09/09 05h40	348	09/09 06h10
17	539	09/09 09h30	345	09/09 10h20	446	09/09 09h00	307	08/09 21h30
18	1673	09/09 04h30	1392	09/09 05h00	1291	09/09 04h00	1175	09/09 04h20

Table 5. Sensitivity studies using the CVN model. The reference simulations were performed with the distributed soils and a Manning roughness coefficient of 0.05. US 528, 535, 607 belong to MARINE textural class 4 (loam); US 430 and 529 to class 7 (clay loam); and US 425 to class 11 (clay).

Simulation	Initial water storage deficit (mm)	Maximum peak discharge ($\text{m}^3 \text{s}^{-1}$)	Time of peak (UT) 2002/09/09	Total runoff (mm)	Runoff coefficient
Grabieux #15 (rainfall 404 mm)					
Reference simulation	80	522	05h40	323	0.80
Uniform soil US 430	30	548	05h40	372	0.92
Uniform soil US 528	134	486	05h40	269	0.66
Uniform soil US 535	64	548	05h40	338	0.84
Uniform soil US 528	80	548	05h40	320	0.79
Courme #17 (rainfall 359 mm)					
Reference simulation	131	539	09h30	271	0.75
Uniform soil US 528	123	570	09h20	232	0.64
Uniform soil US 529	172	514	09h30	255	0.71
Uniform soil US 430	52	573	09h20	299	0.83
Uniform soil US 529	131	514	09h30	255	0.71
Uniform soil US 529 vertically homogeneous	166	513	09h30	247	0.69
Saumane #18 (rainfall 208 mm)					
Observation	-	770	06h00	85	0.41
Reference simulation	36	1673	04h40	175	0.84
Uniform soil US 607	48	1696	04h30	156	0.75
Vertically homogeneous soil	38	1674	04h30	170	0.82

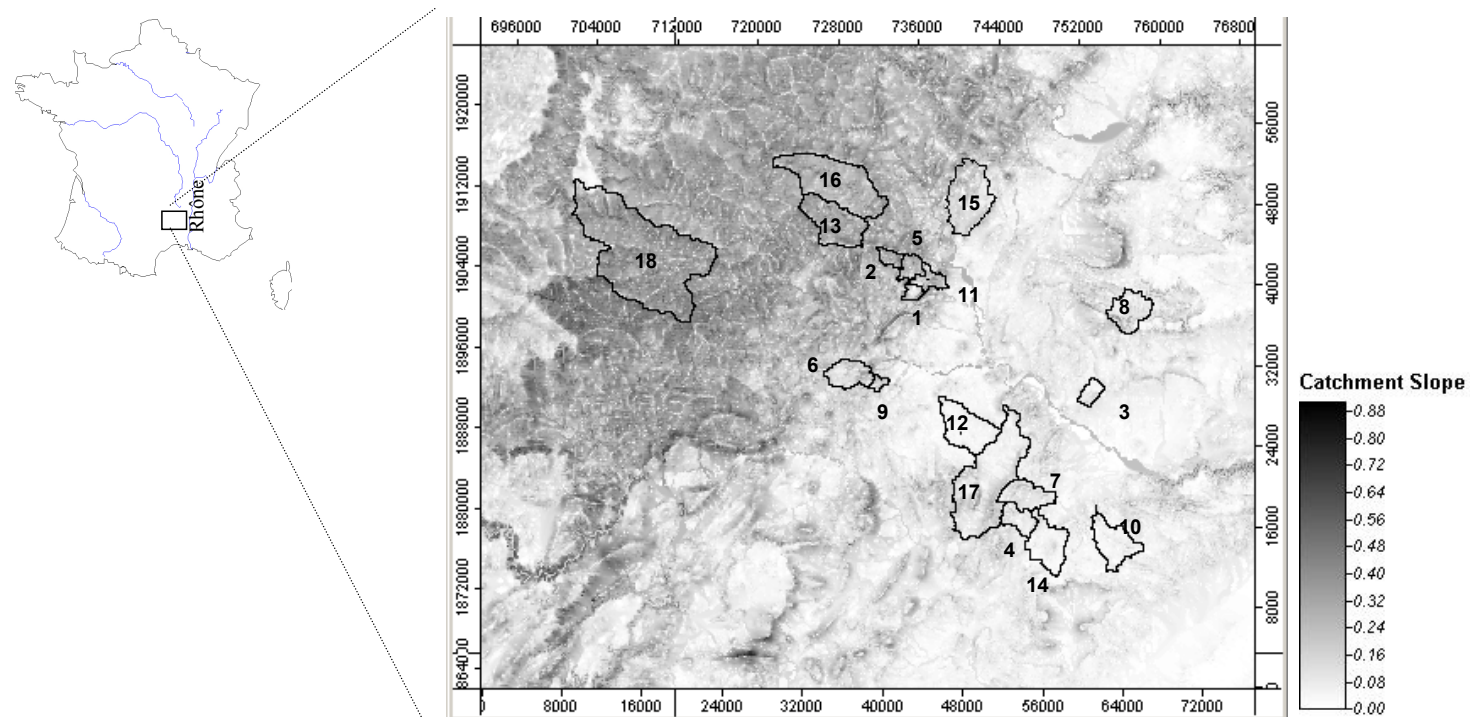


Figure 1.

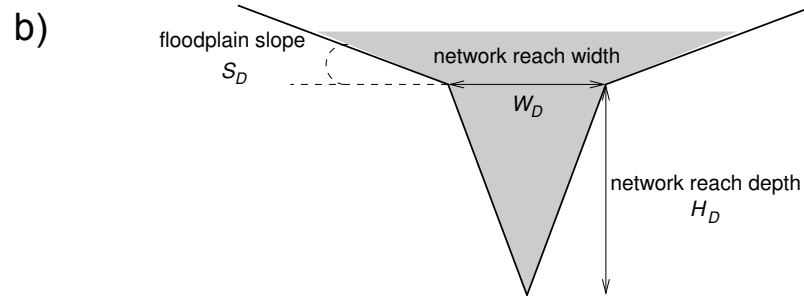
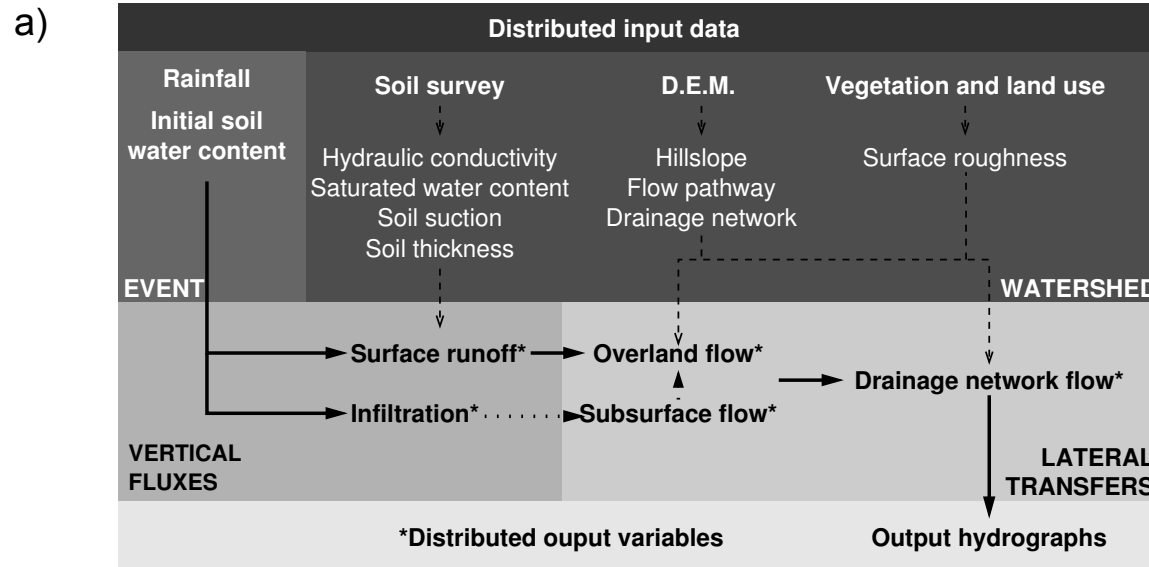


Figure 2

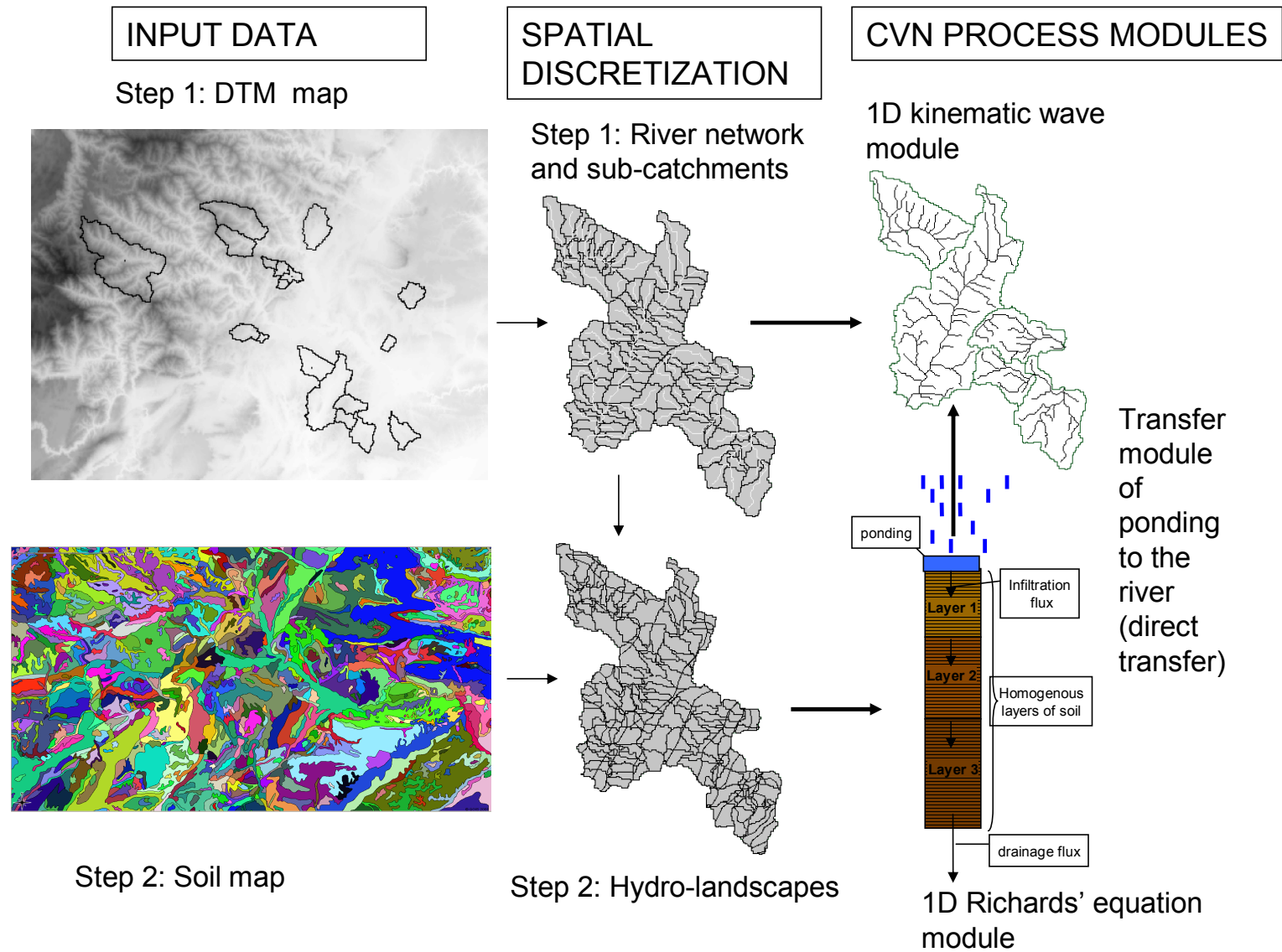


Figure 3

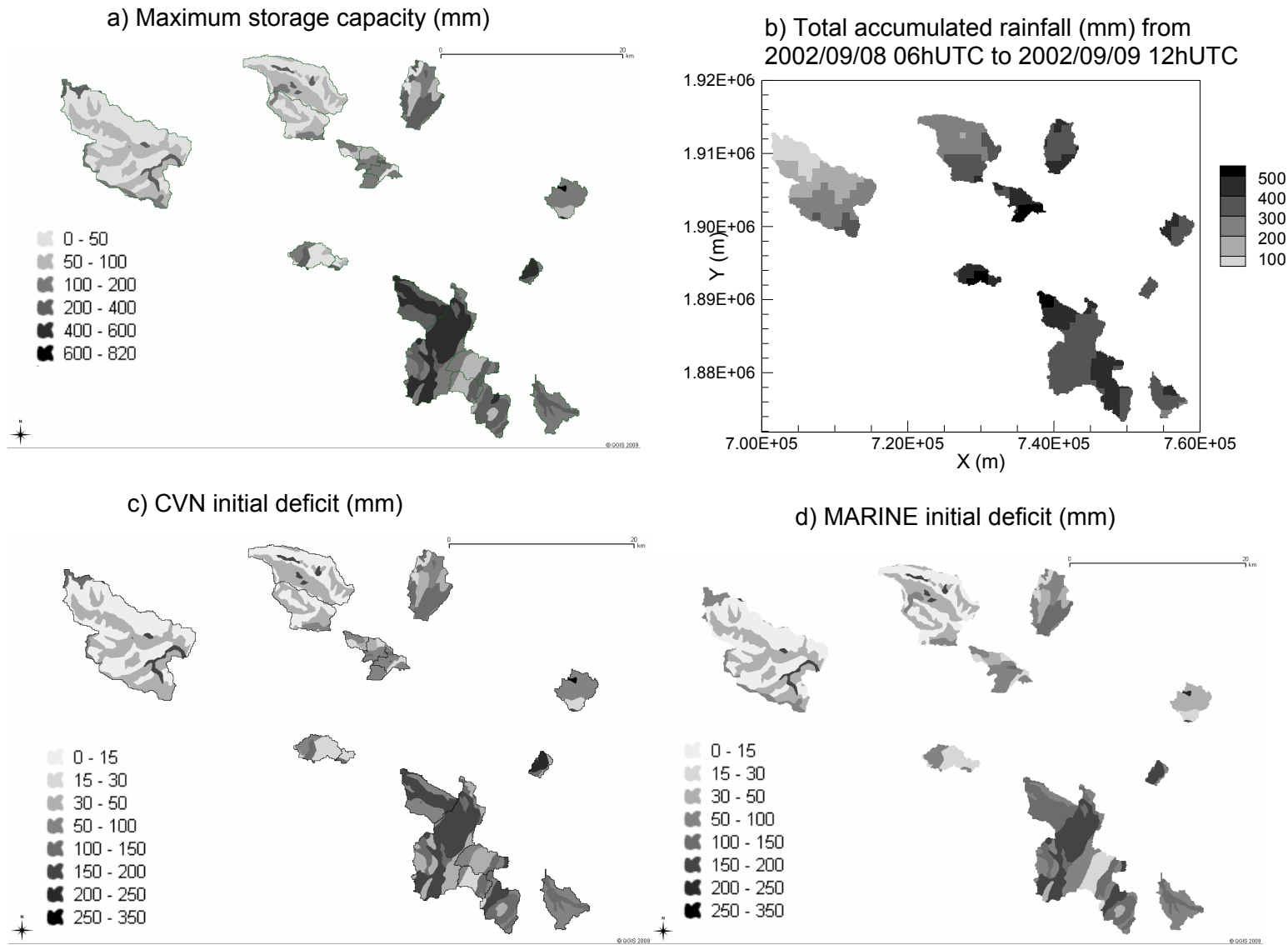


Figure 4

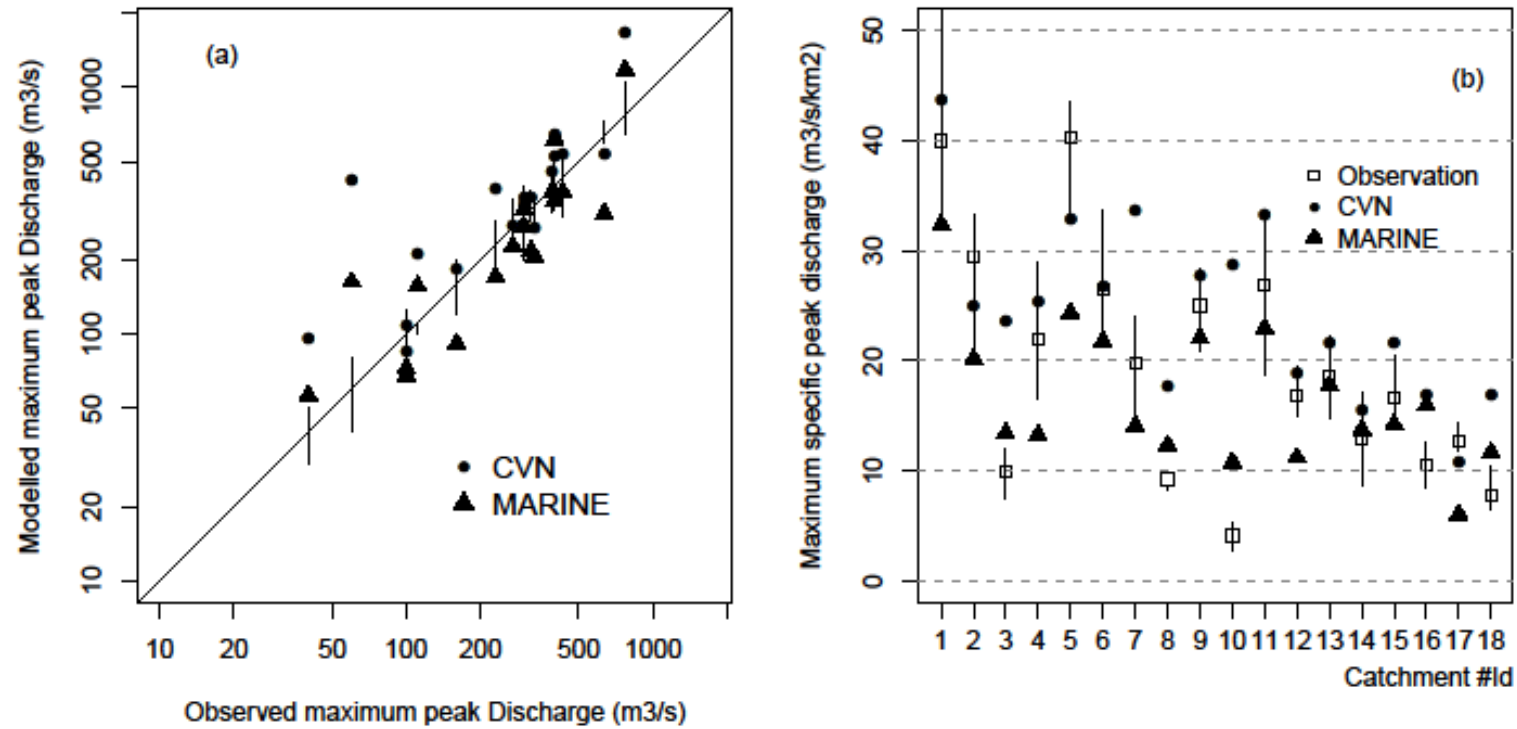


Figure 5

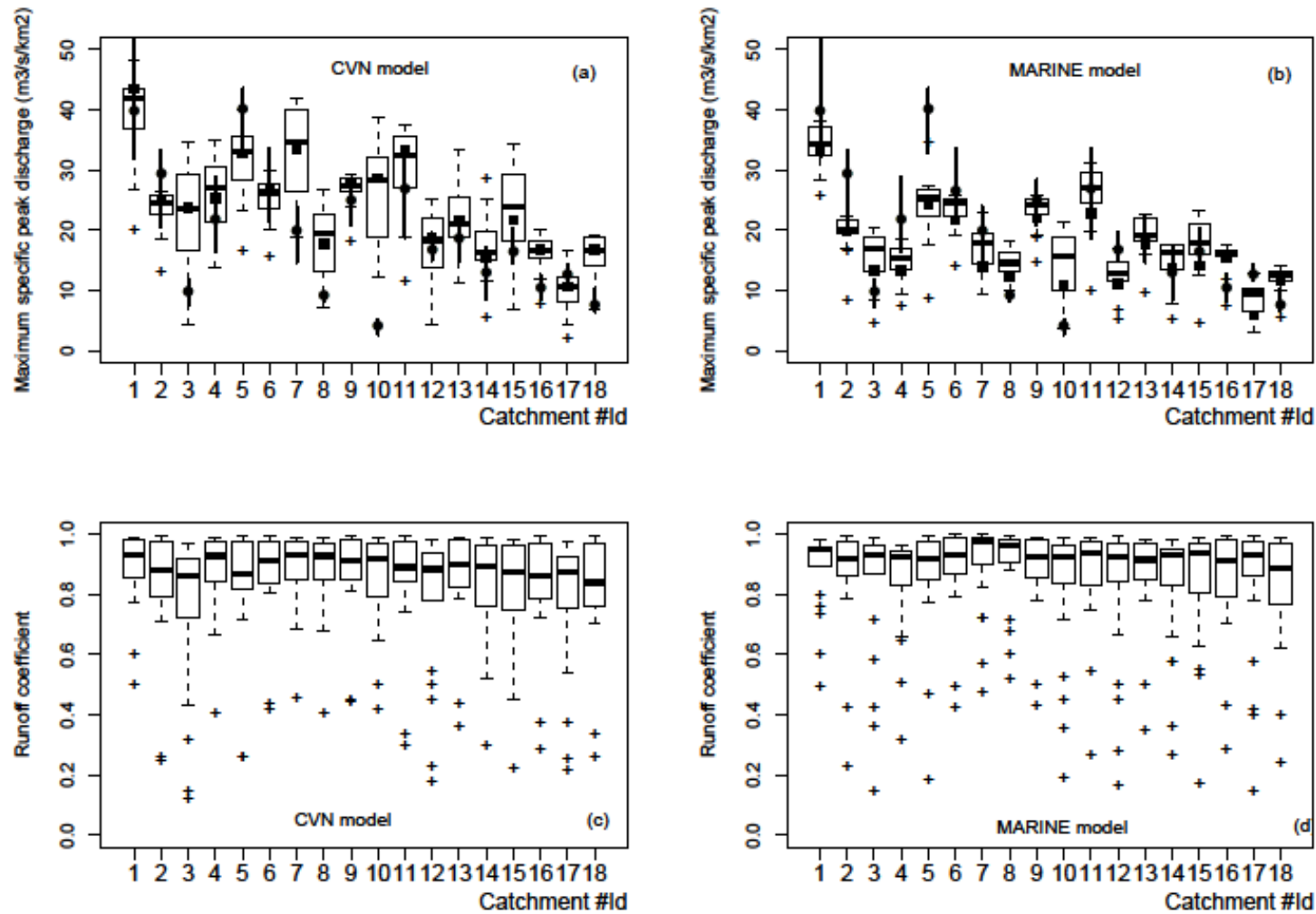


Figure 6

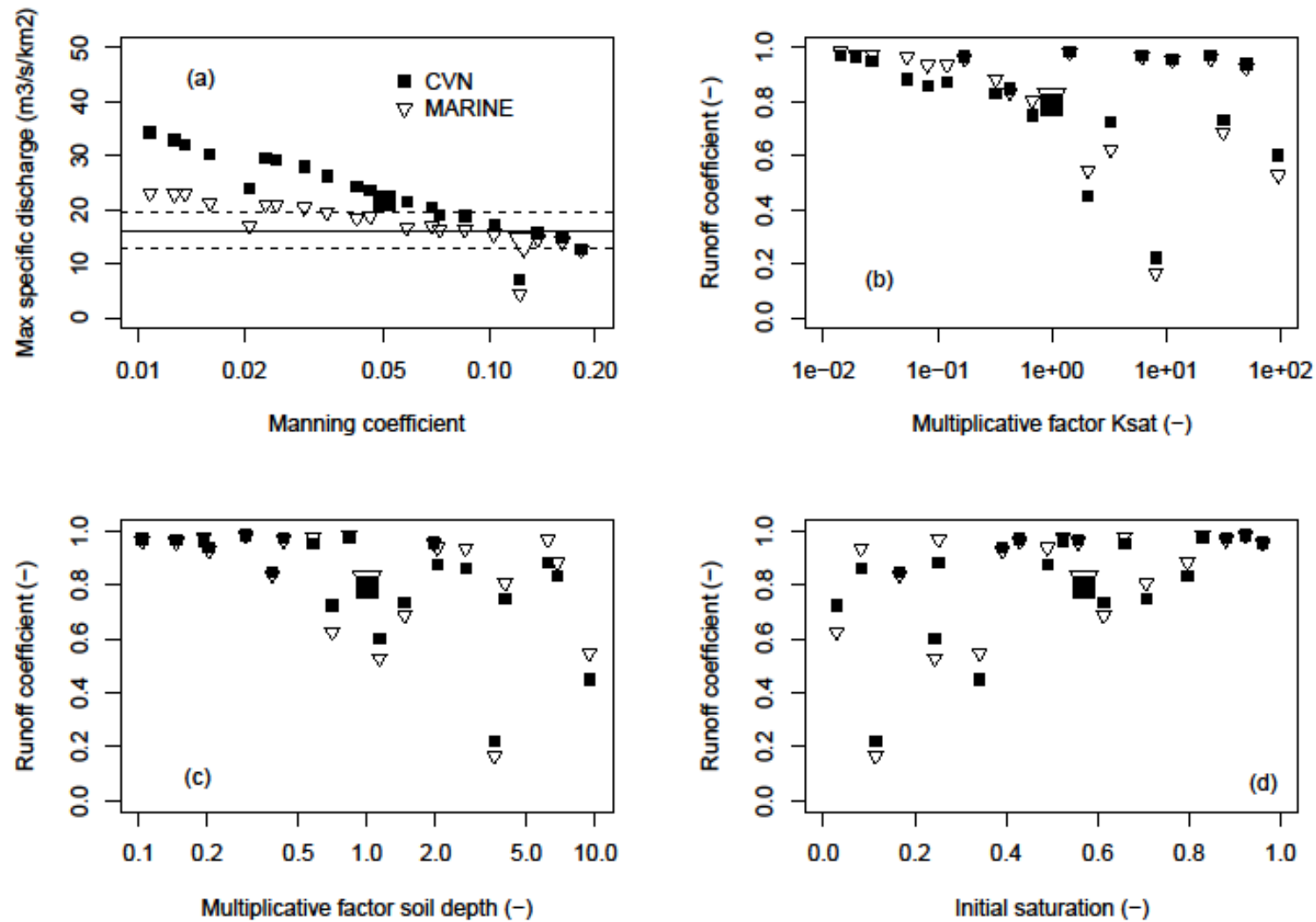


Figure 7

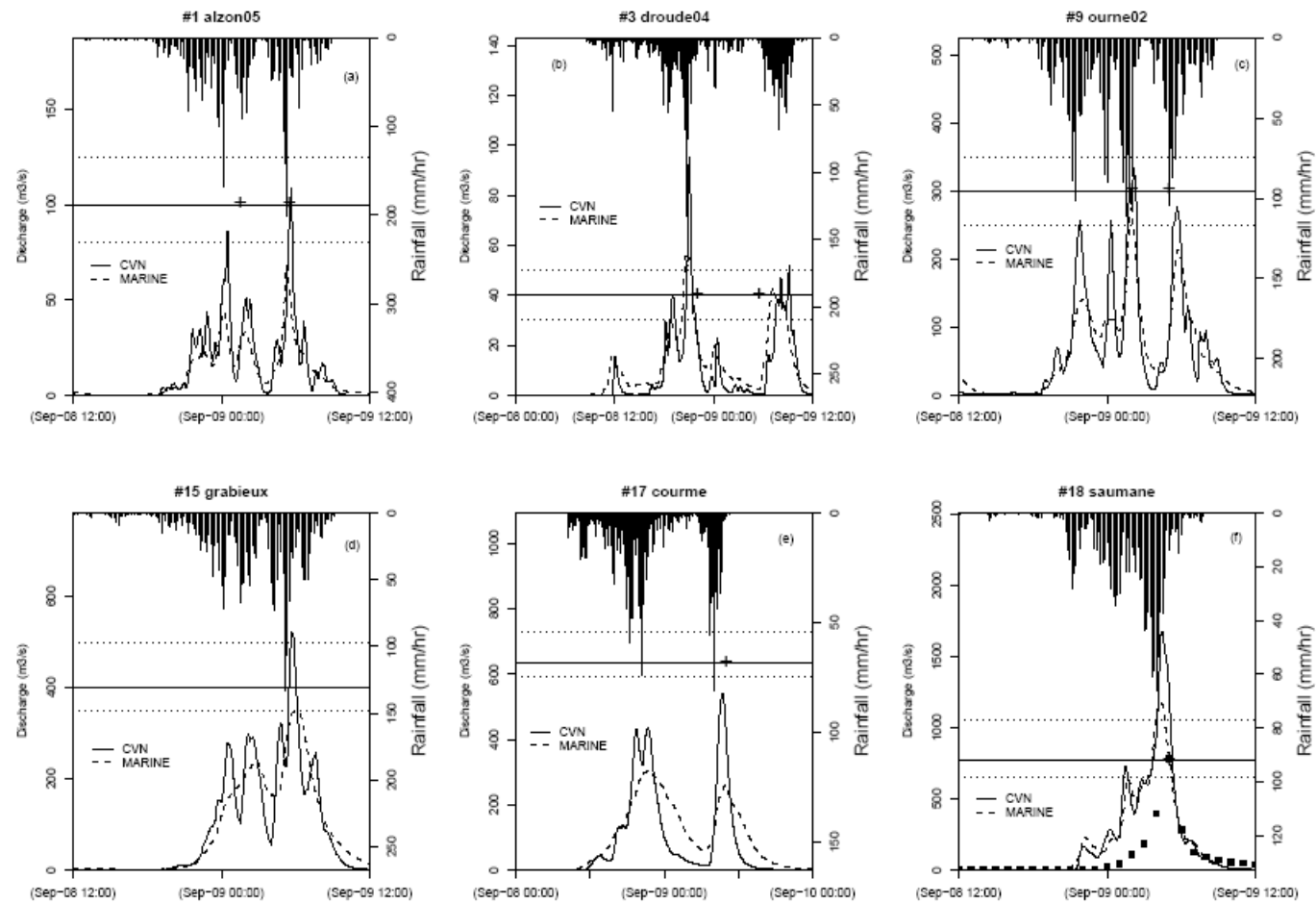


Figure 8

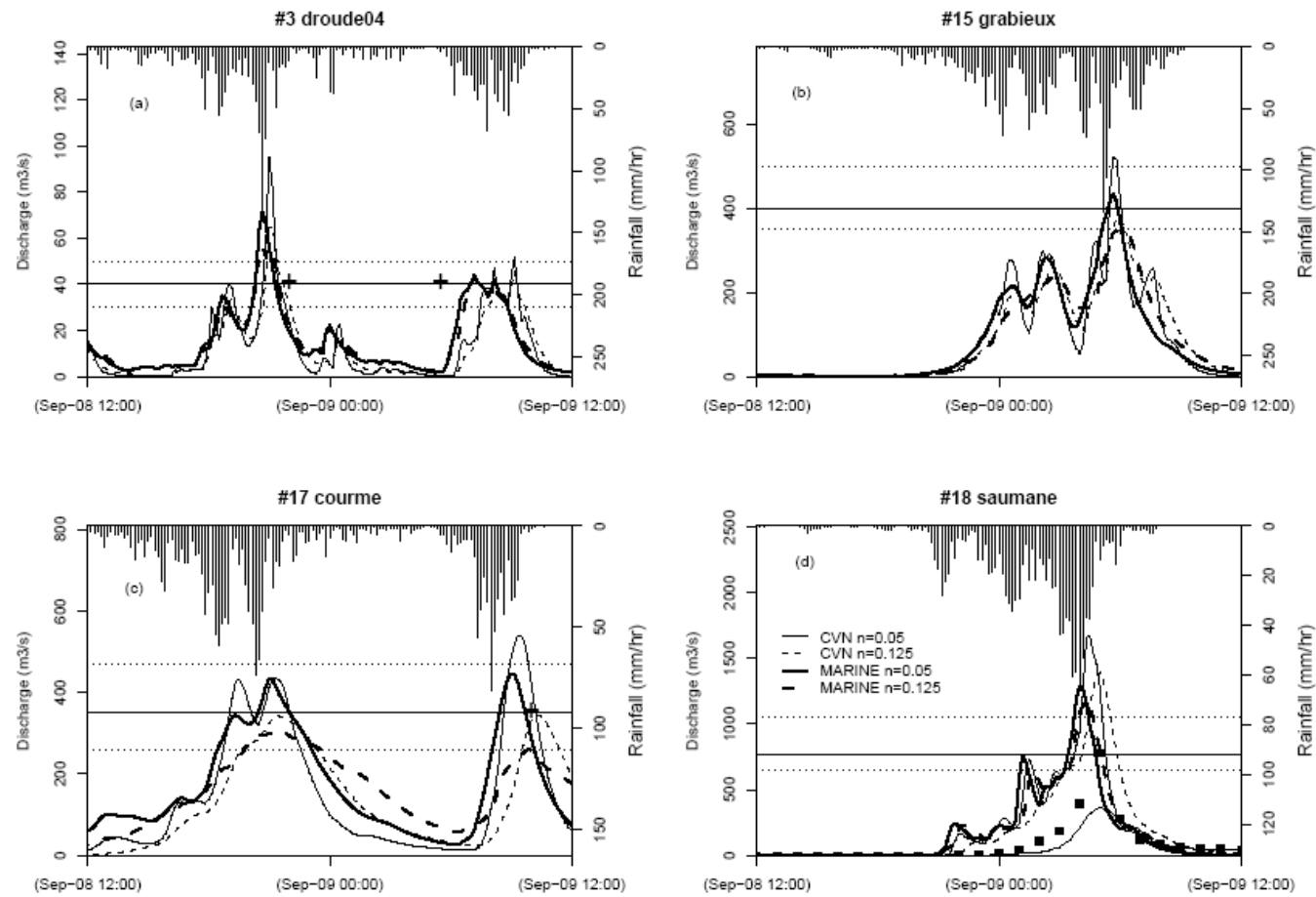


Figure 9

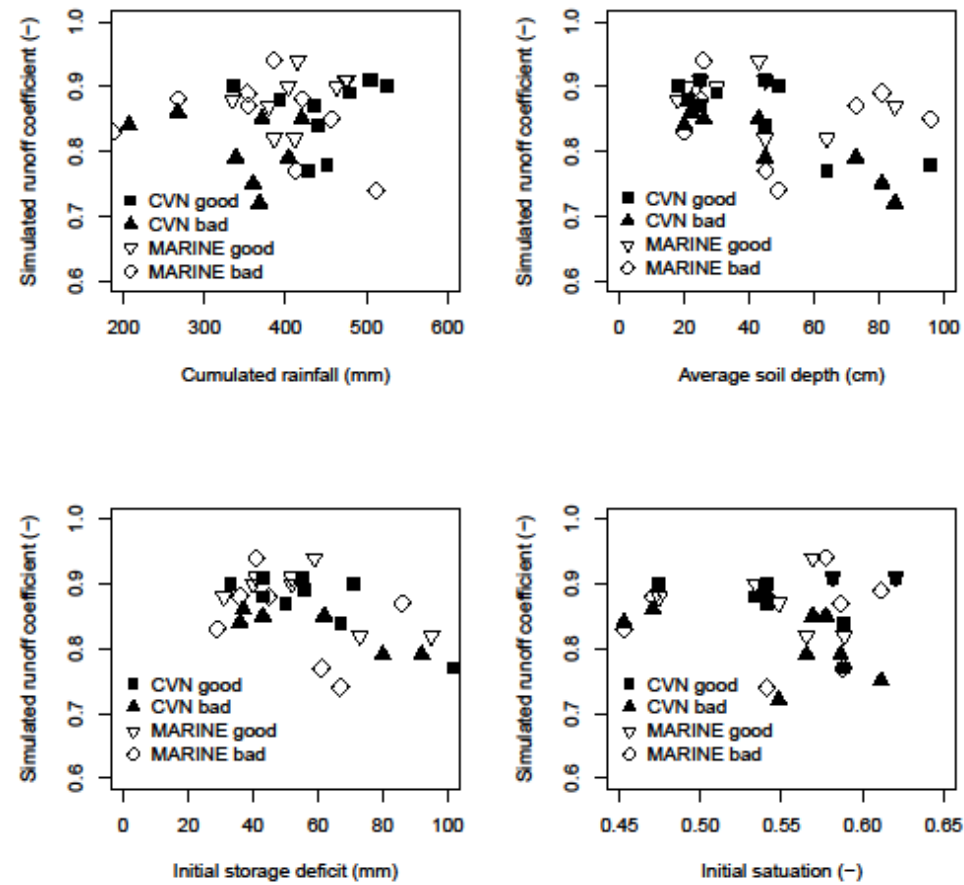


Figure 10

Figure 10. Spatial distribution and time evolution of the saturation state of each catchment with a) CVN model b) MARINE model

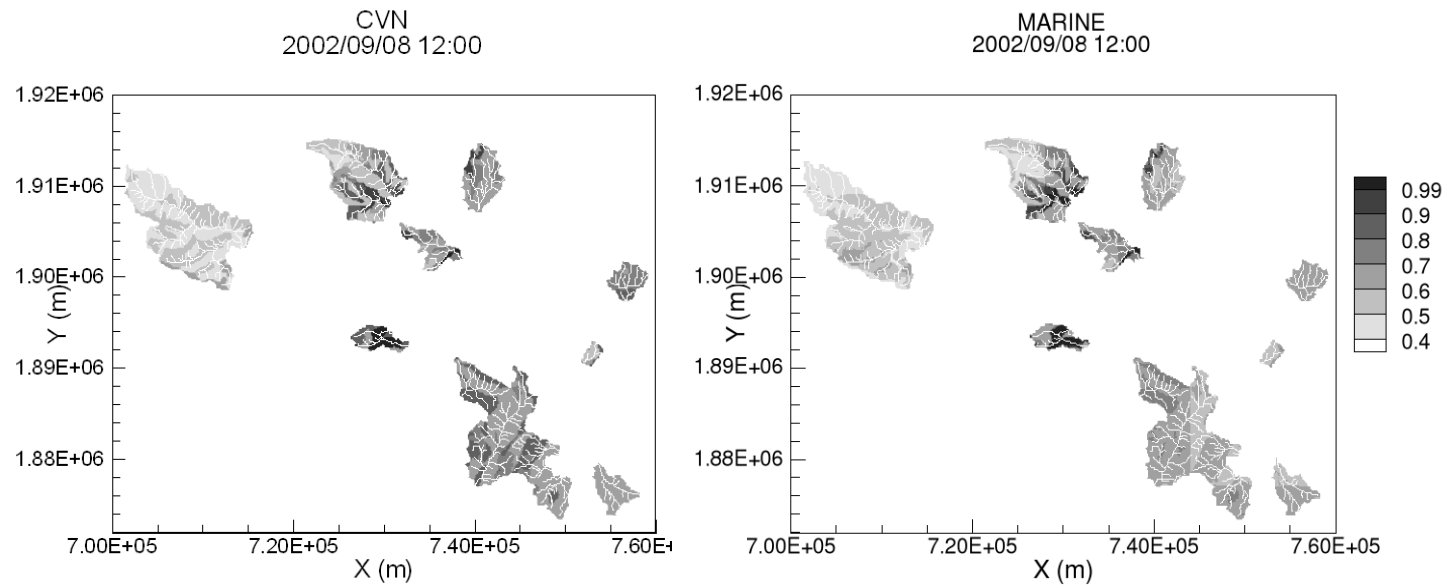


Figure 11

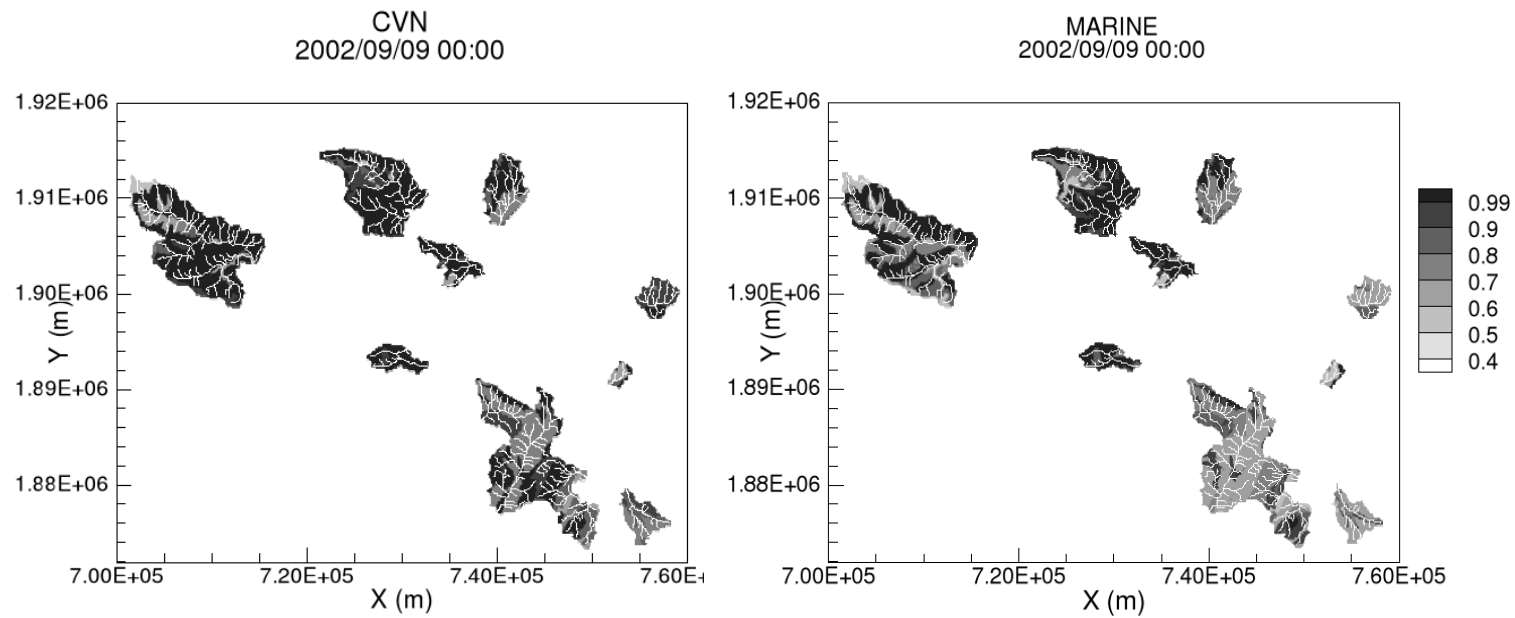


Figure 11 Continued

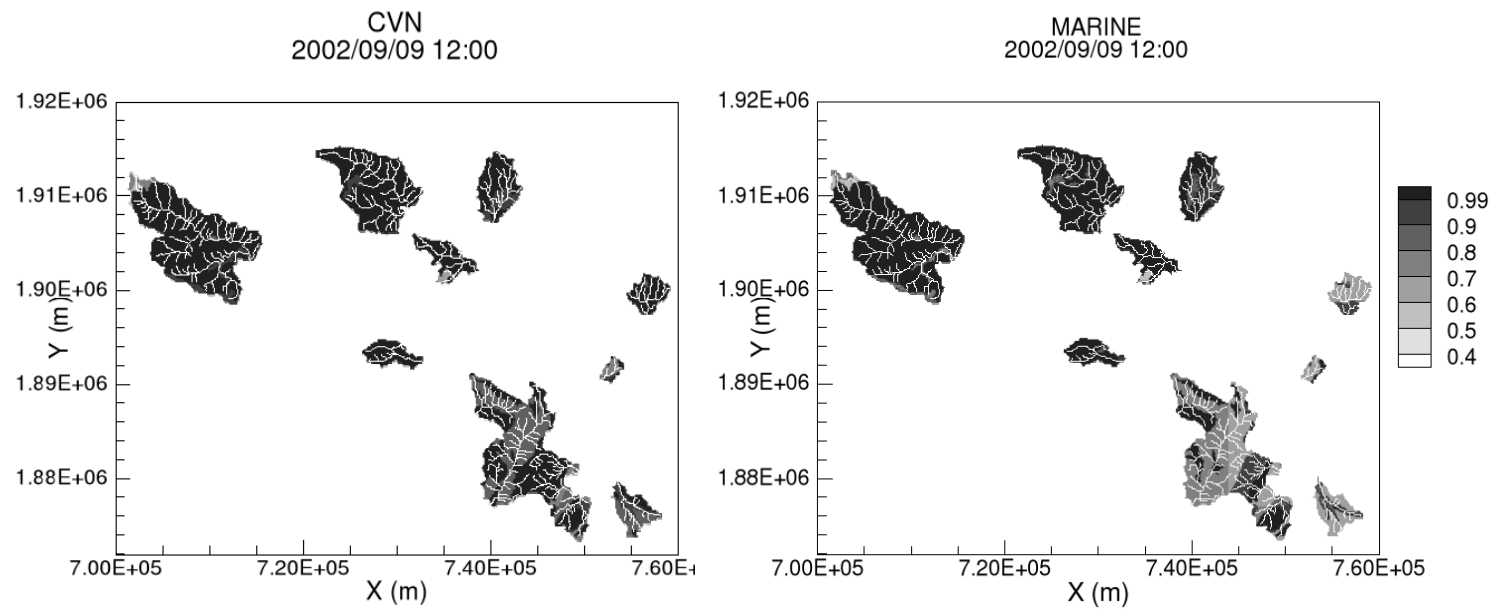


Figure 11 Continued

Figure 9. Time evolution of the mean saturation state of all the studied area

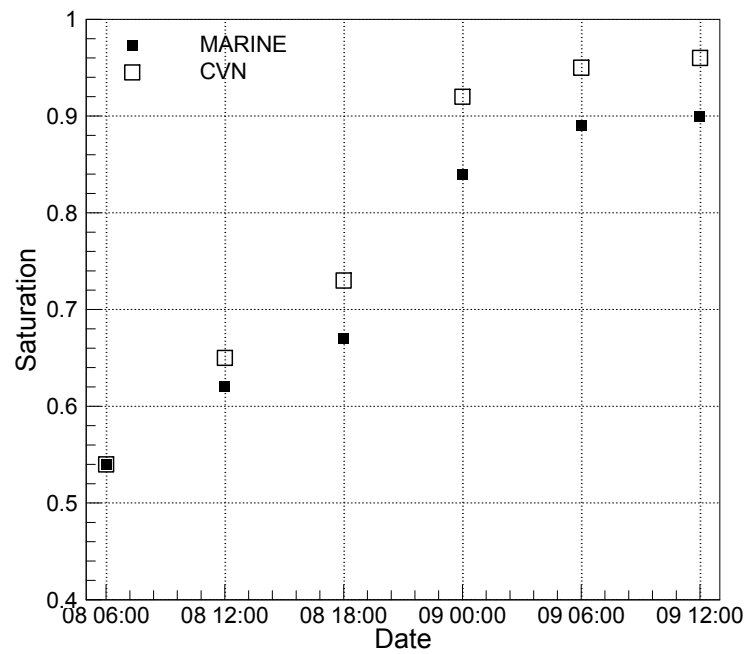


Figure 12

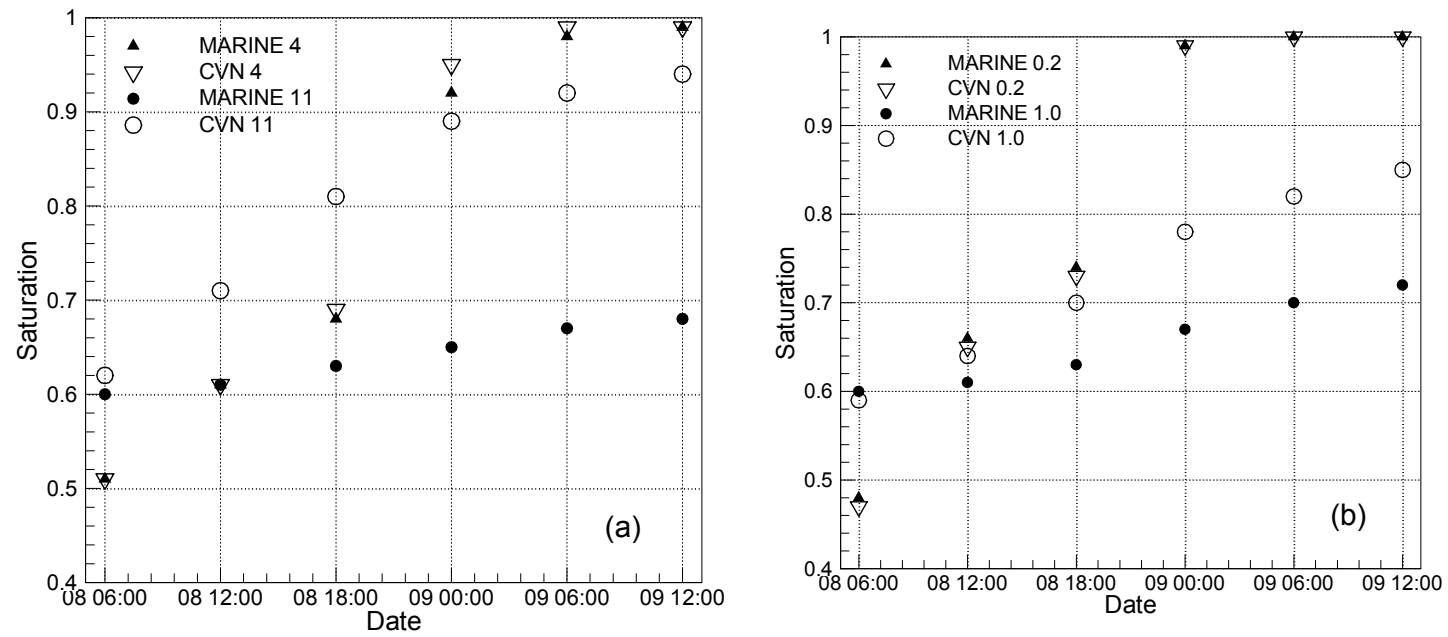


Figure 13

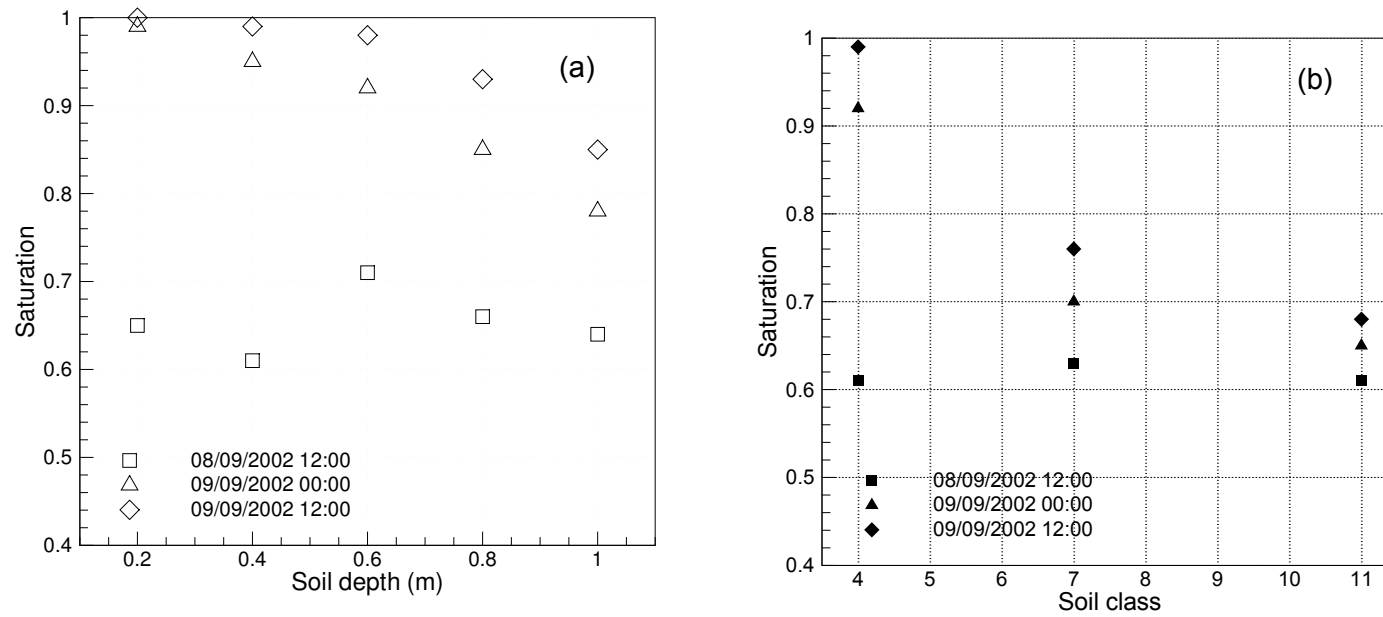


Figure 14

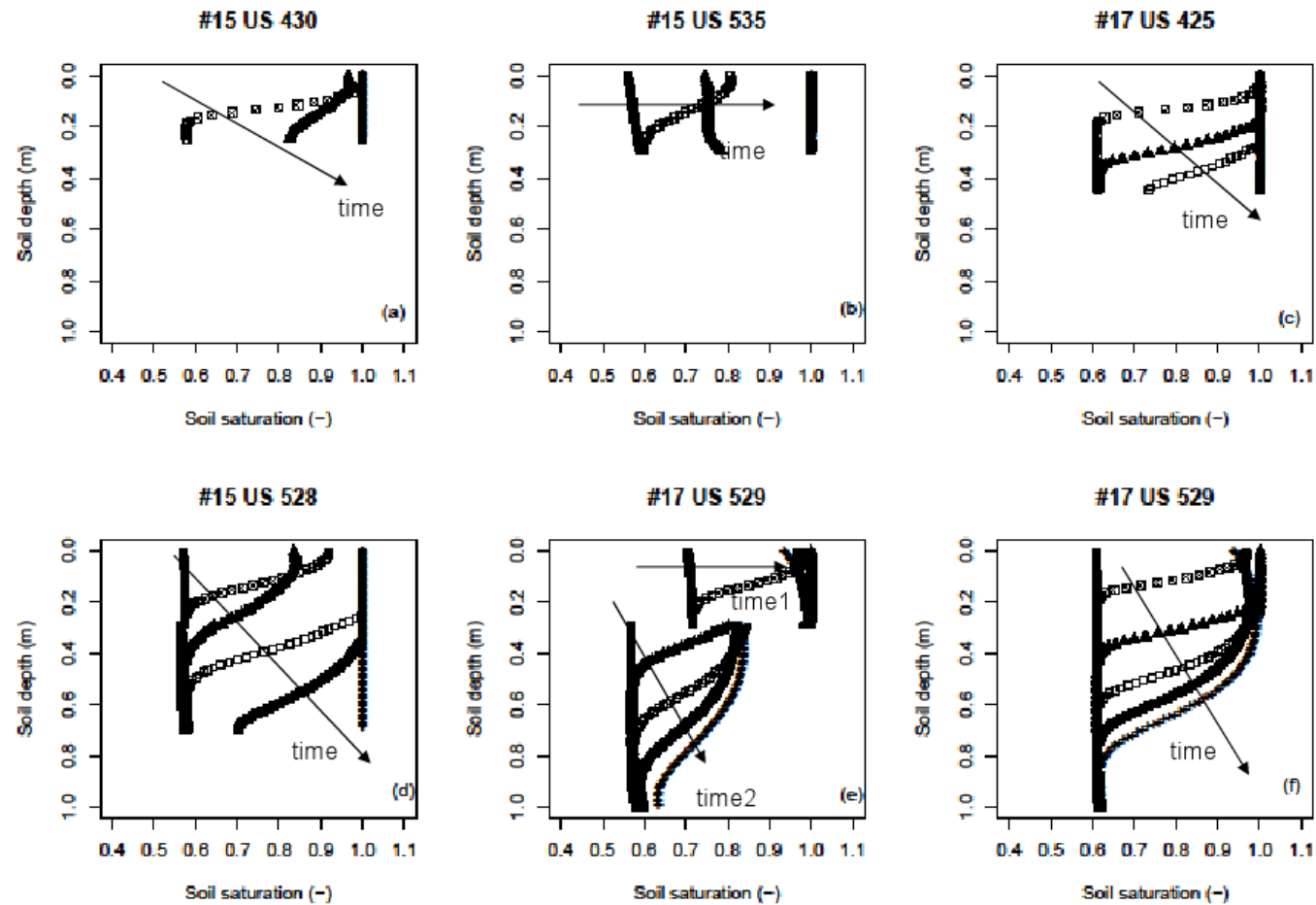


Figure 15

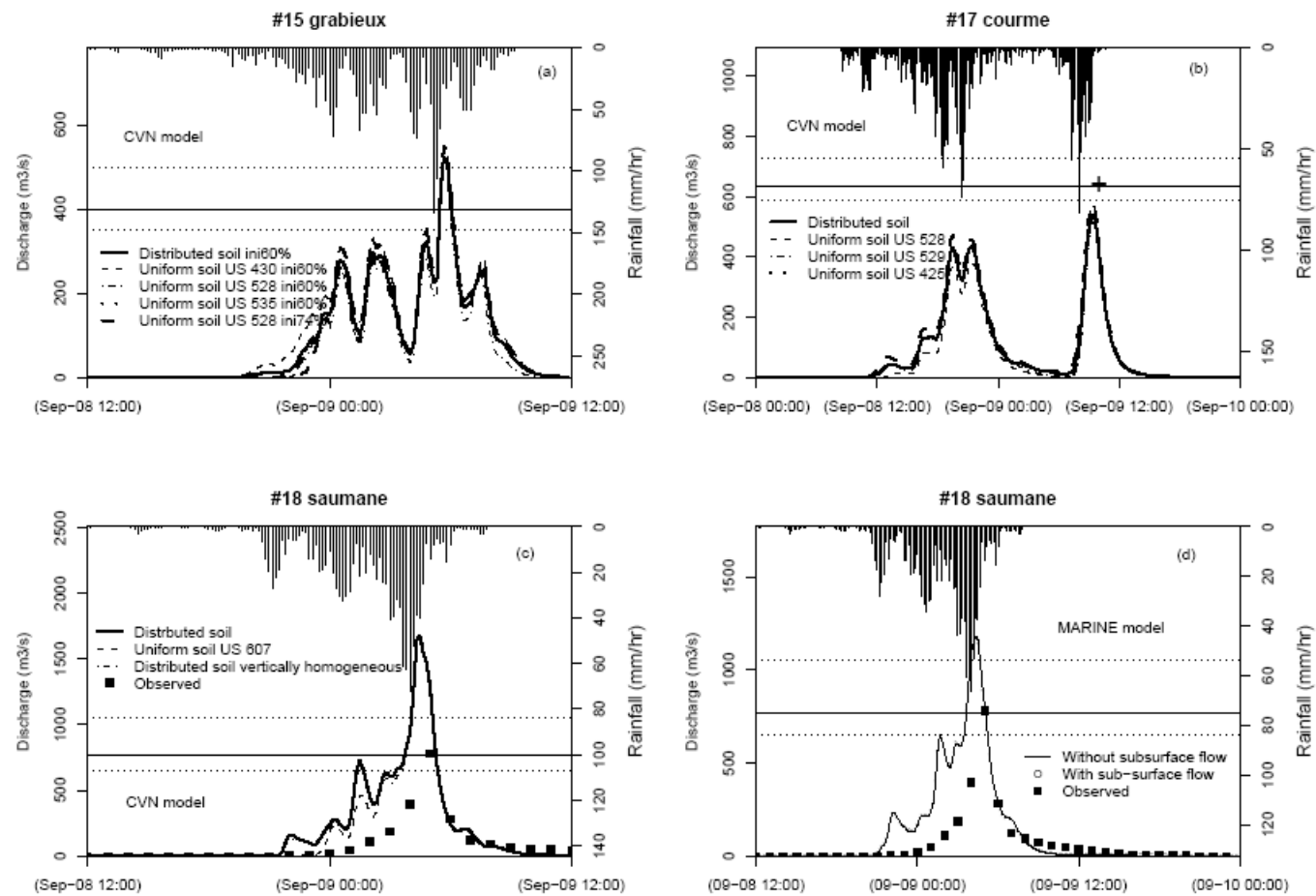


Figure 16

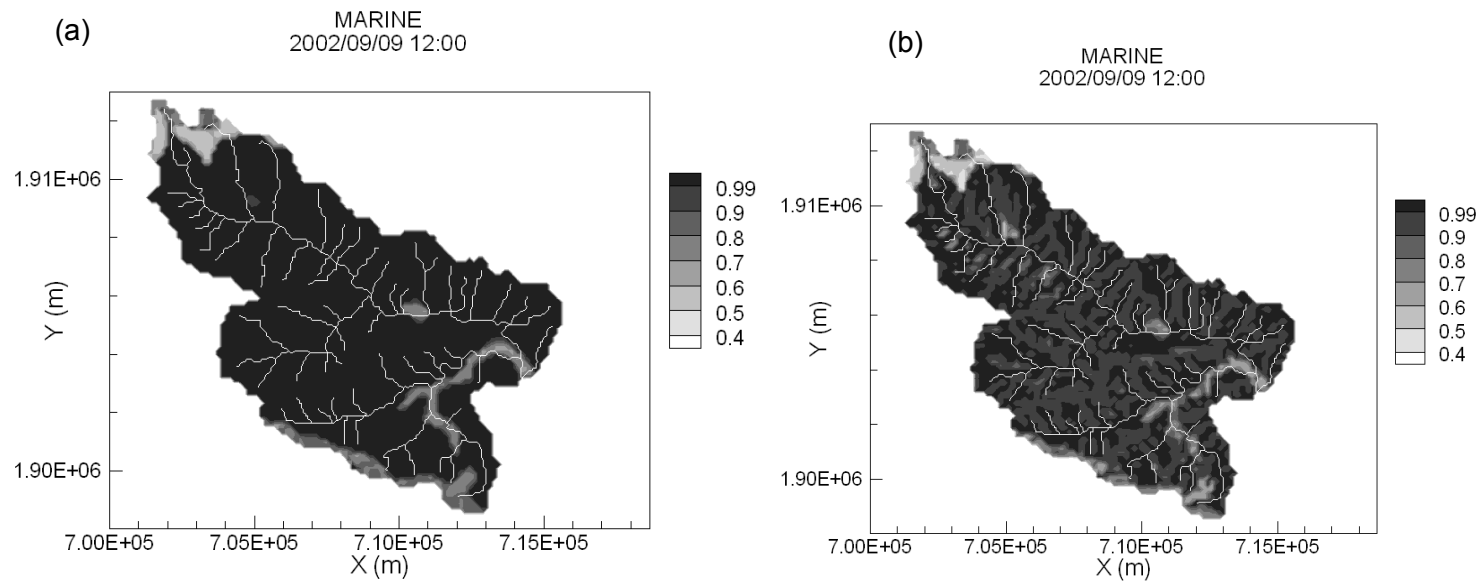


Figure 17